



Over de zoutcrisis in de Middellandse Zee tijdens
het Messinien (2)

Voor de werkgroep Algemene Geologie Kring Rijnland

Tom Kleijn 15 december 2015

Inhoud

- Inleiding
- Het Messinien en MSC
- Watercirculatie in MZ.
- De zeestraten tussen de AO en de MZ
- De aanloop naar de zoutcrisis en de eerste afzettingen van de zoutcrisis

DISTRIBUTION OF EMERGED LAND AND SEA 6 MILLION YEARS AGO





S

MSC onset

Perales section
(Sierra et al., 2001)

PRIMARY LOWER GYPSUM
(Yesares Formation)

PLG2

PLG1

UA34

UA33

UA32

UA31

UA30

UA34

UA33

UA32

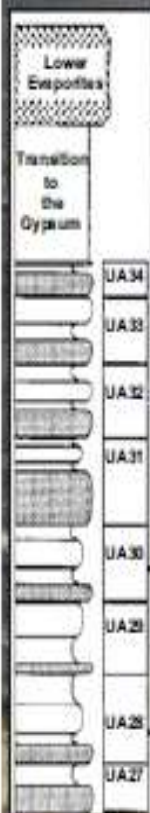
UA31

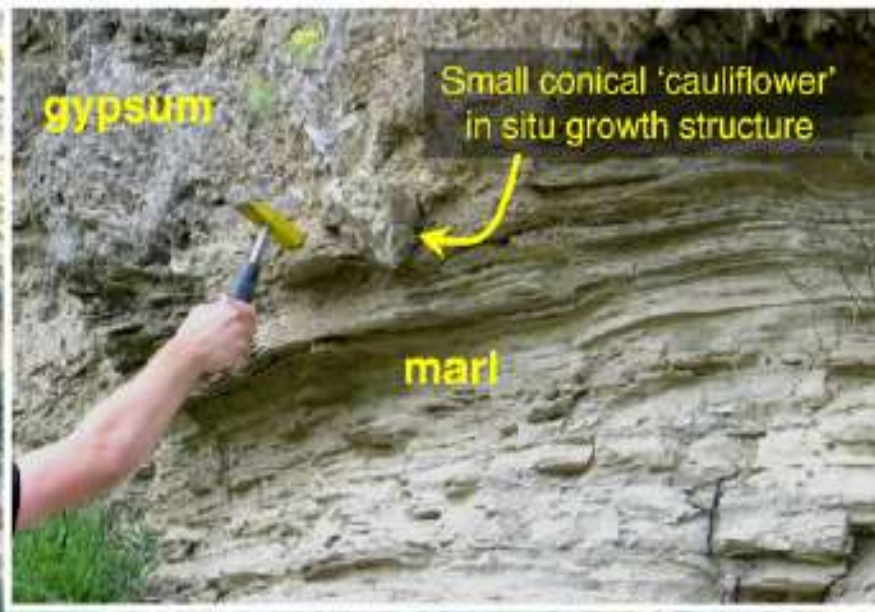
UA30

UA29

UA28

PRE-EVAPORITIC
(Abad Formation)





Yesares Formation, Sorbas Basin, S. Spain. Scale: horizontal ~15m; vertical ~5m



Wanneer?

Period	Epoch	Stage		
Neogene	Holocene			
	Pleistocene	L		
		M		
		E		
	Pliocene	L	Gelasian	← GSSP 1985 (1.806 Ma)
		M	Piacenzian	← GSSP 1996 (2.588 Ma)
		E	Zanclean	← GSSP 1997 (3.600 Ma)
			Messinian	← GSSP 2000 (5.333 Ma)
	Miocene	L	Tortonian	← GSSP 2000 (7.246 Ma)
			Serravallian	← GSSP 2003 (11.608 Ma)
		M	Langhian	← GSSP 2007 (13.82 Ma)
			Burdigalian	
		E	Aquitanian	
				↑ Working groups ↓

De “Golden Spikes” voor het Messinien

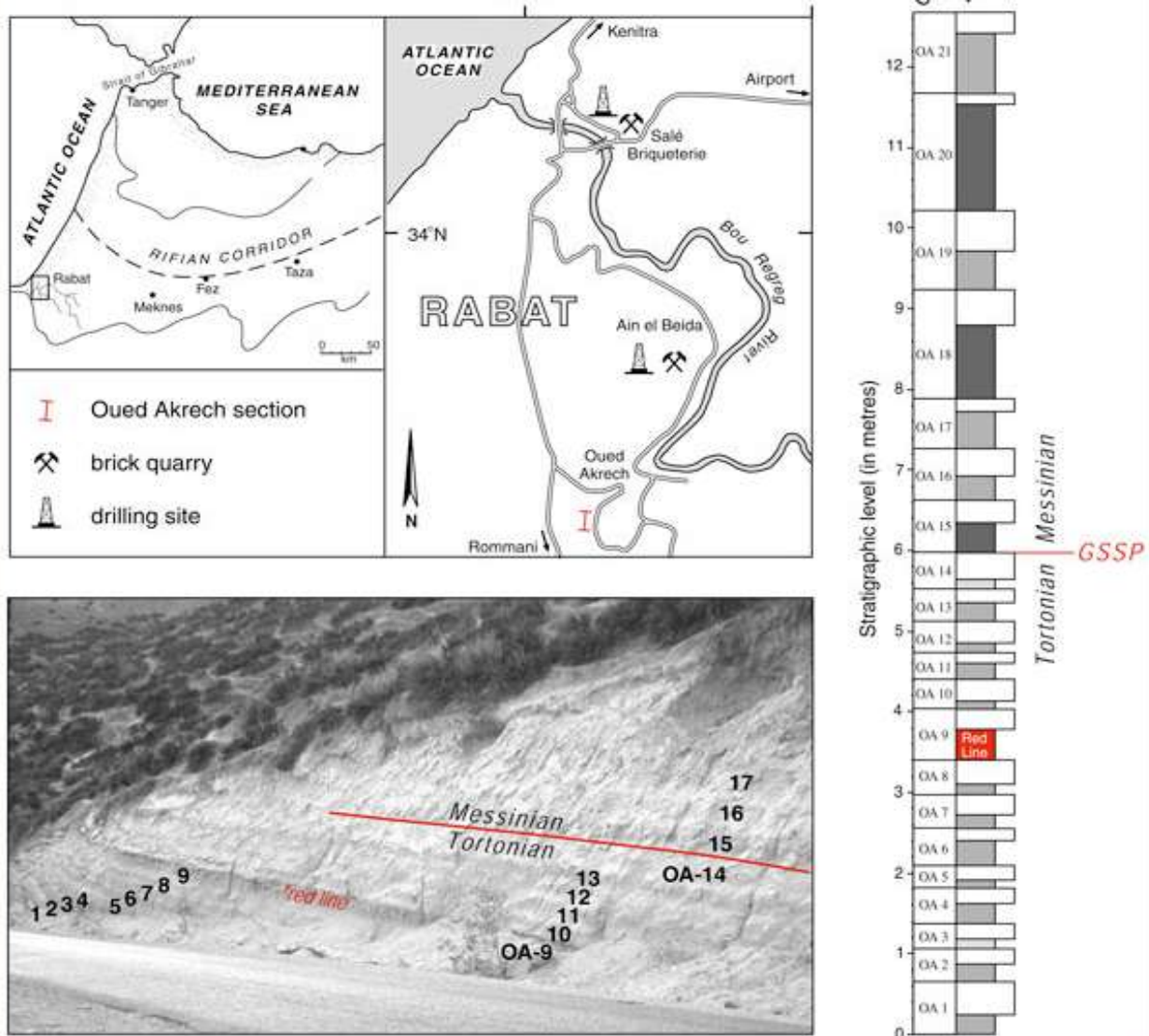
GSSP's

Map of GSSP's



Stage	Section	Location
base Pleistocene	Vrica, near Crotona, Italy	39°05' N, 17°07' E
base Gelasian	Mt S. Nicola, near Gela, Sicily, Italy	37°04' N, 14°15' E
base Piacenzian	Punte Piccola, Sicily, Italy	37°17'20" N, 13°29'36" E
base Zanclean	Eraclea Minoa, Sicily, Italy	37°23'30" N, 13°16'50" E
base Messinian	Oued Akrech, Morocco	33°56'13" N, 6°48'45" E
base Neogene	Lemme-Carrosio, northern Italy	44°39'32" N, 8°50'11" E
base Serravallian	Ras il Pellegrin	35°54'56" N, 14°20'03" E

Begin Messinien



Het einde van het Messinien

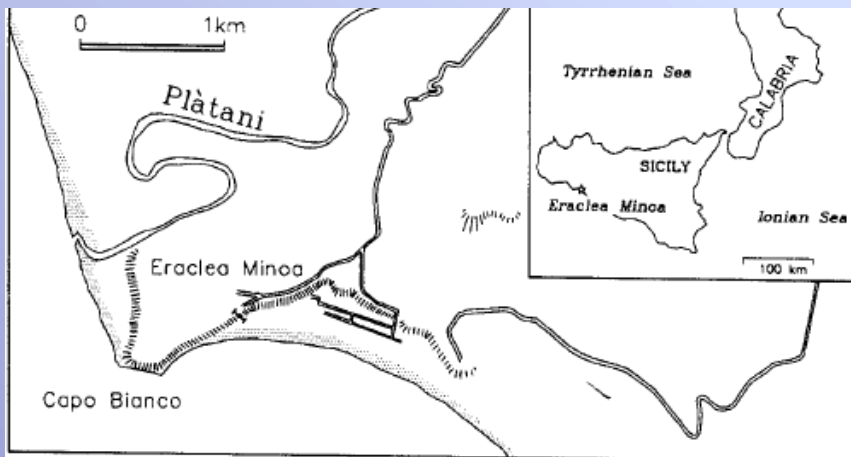
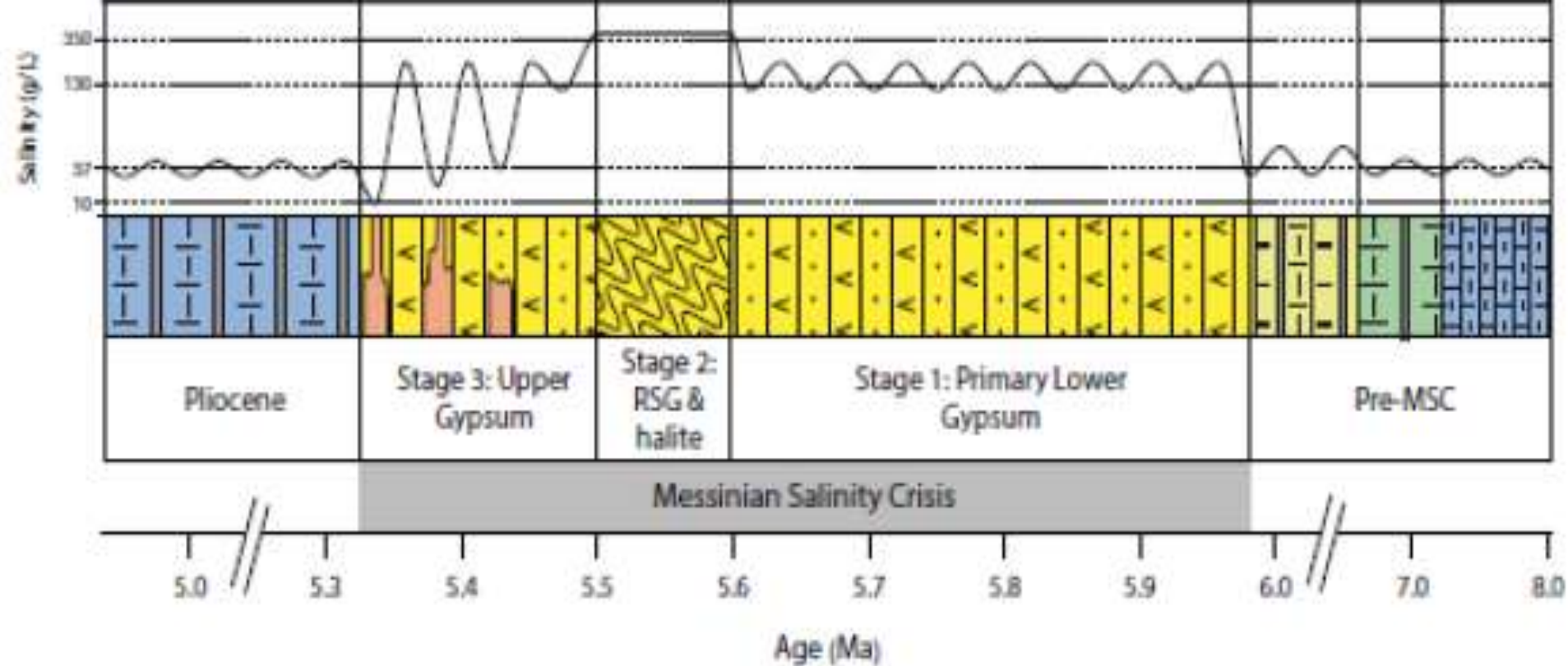


Figure 1 Location of the Eraclea Minosa section.

Eraclea Minoa section, southern coast Sicily, Italy







Indeling Messinien 7,246 Ma

Begin Messinien 7,246

Einde Messinien 5,33

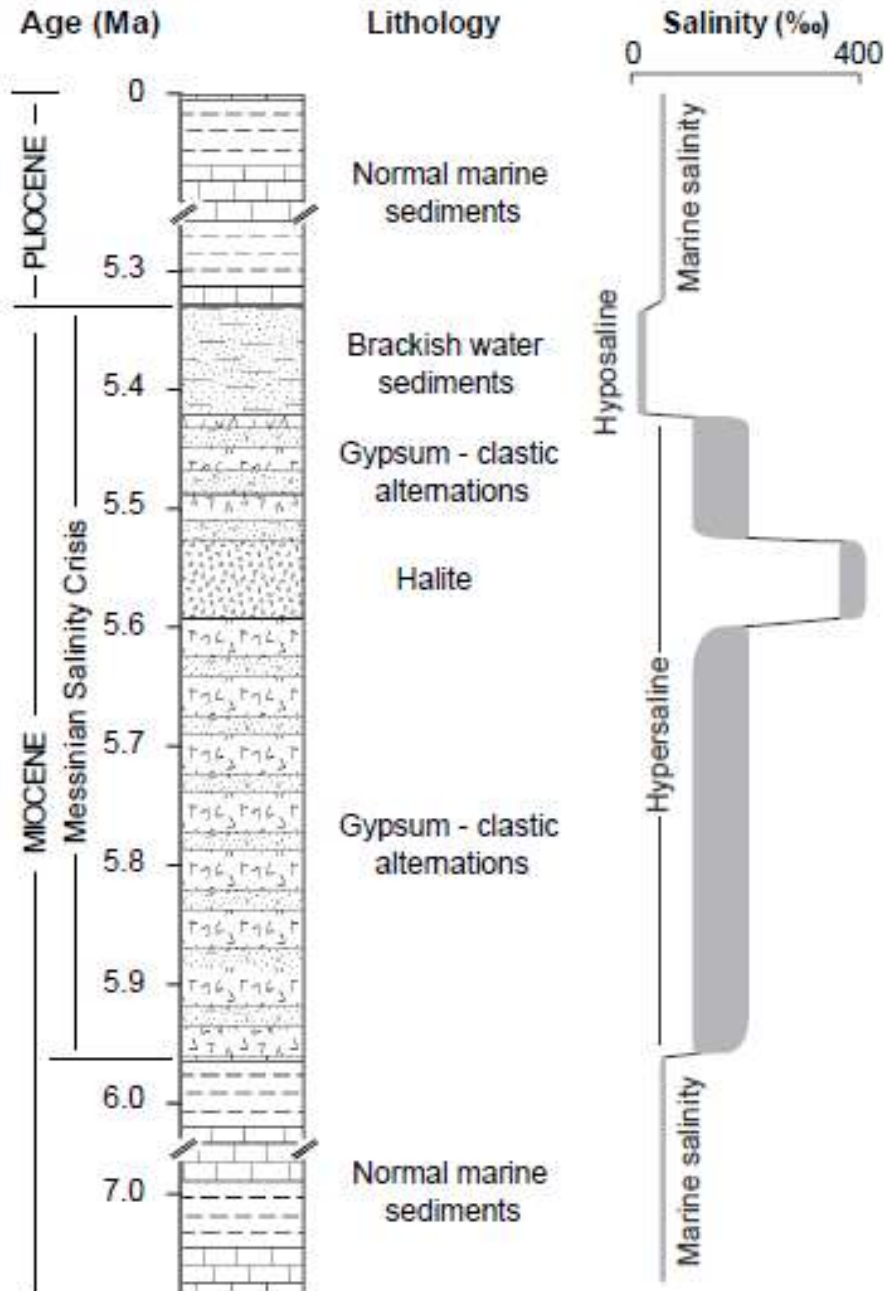
Totale duur Messinien 1,913 milj. jaar

Aanloop naar MSC ruim 1,275 milj. jaar

Begin MSC 5,971Ma

Einde MSC 5,33 Ma

MSC totaal 638.000 jaar



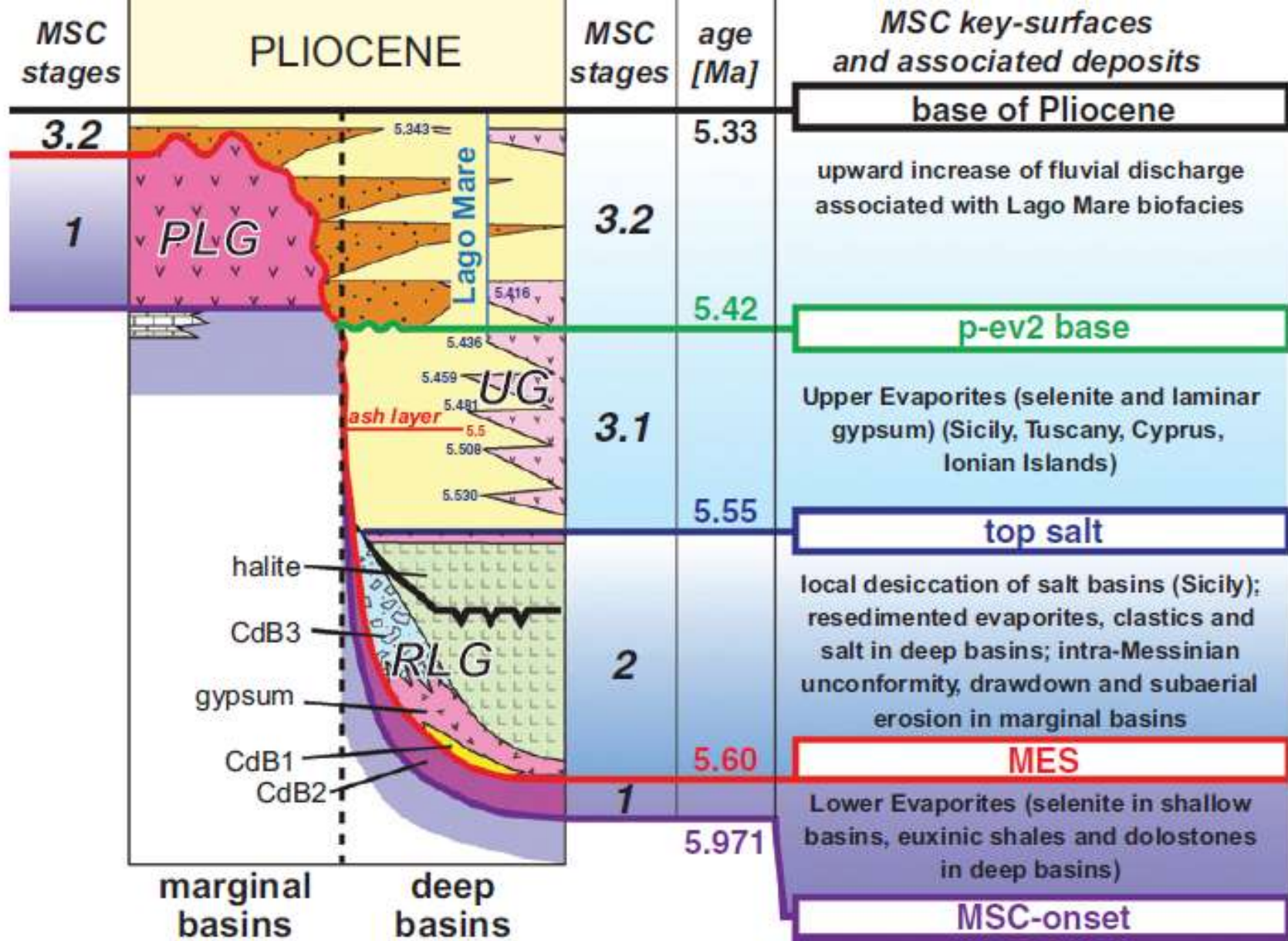
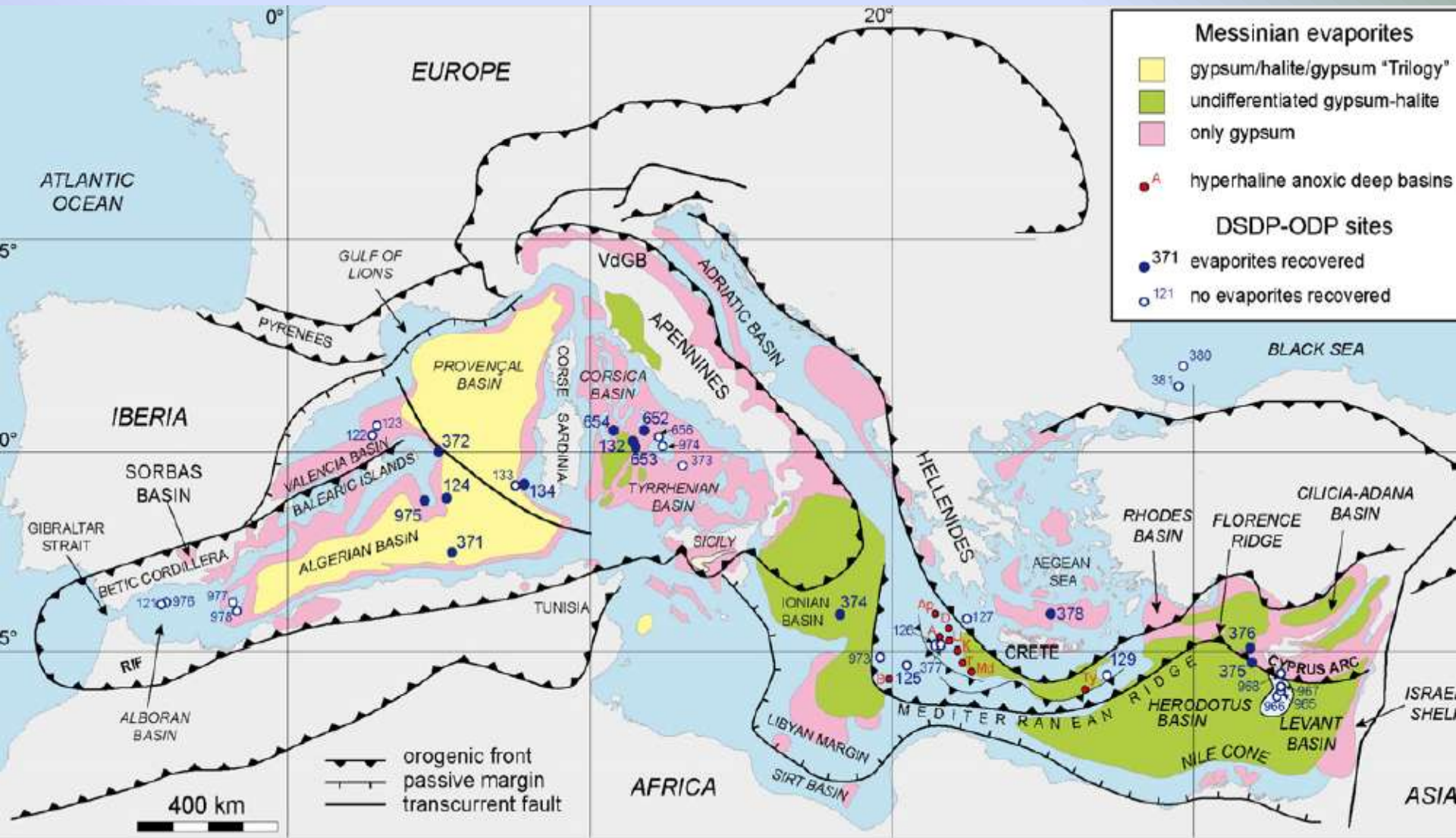
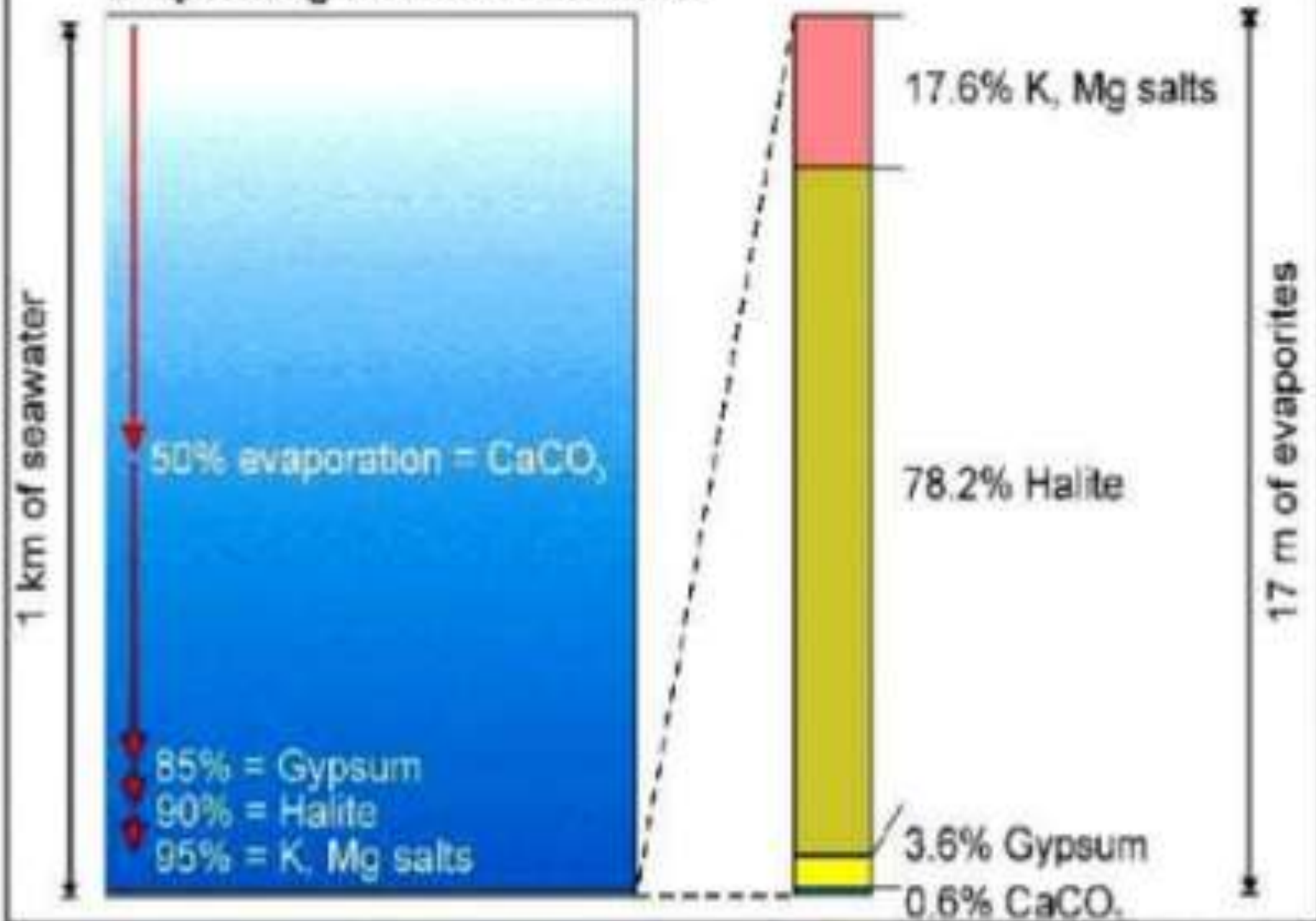


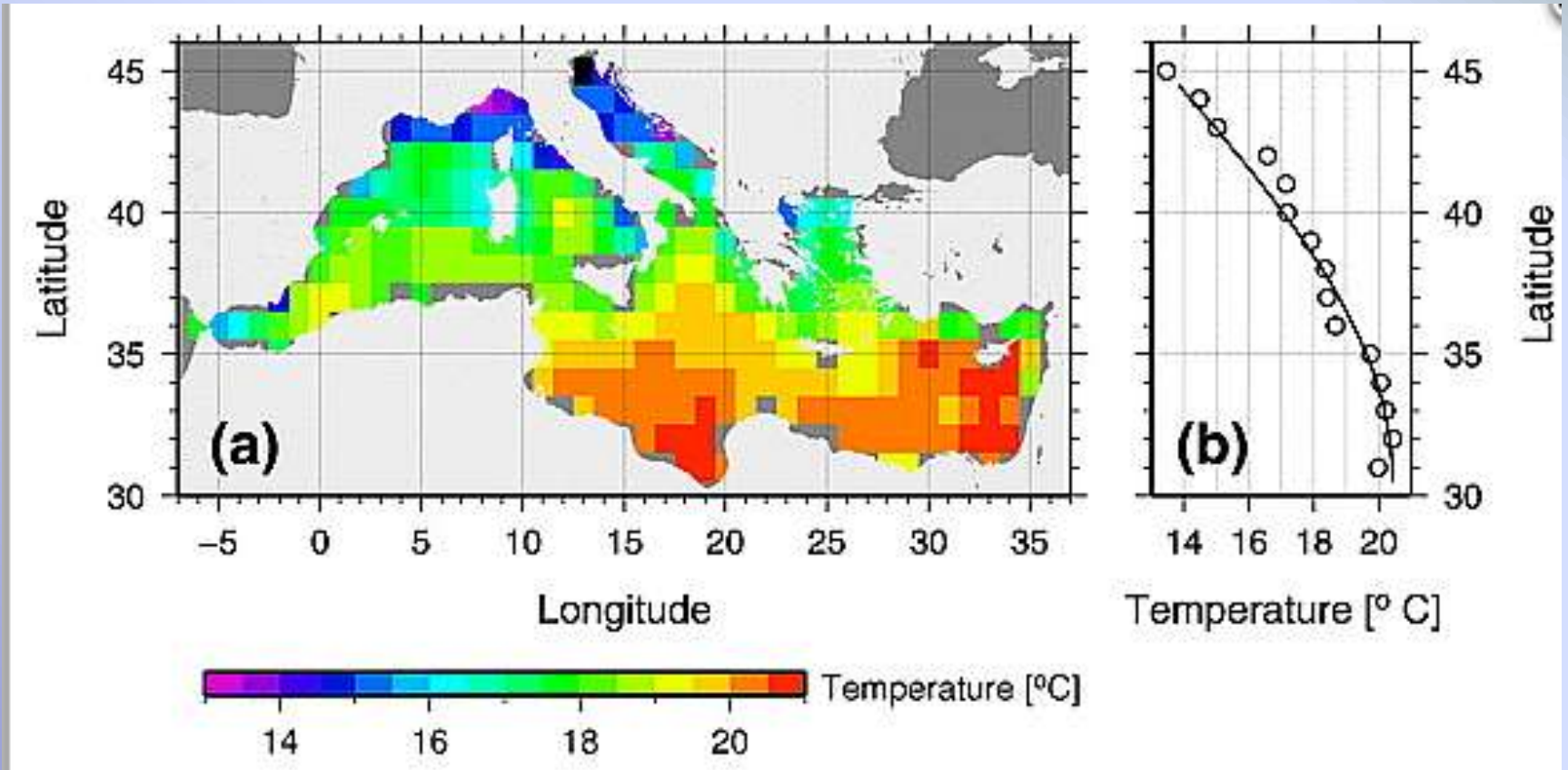
Fig. 1 The CIESM - Commission Internationale pour l'Exploration Scientifique de la mer Méditerranée (2008) Messinian salinity crisis stratigraphic framework (modified after: Roveri *et al.*, 2008a,b; Manzi *et al.*, 2011) showing the 5 key-surfaces



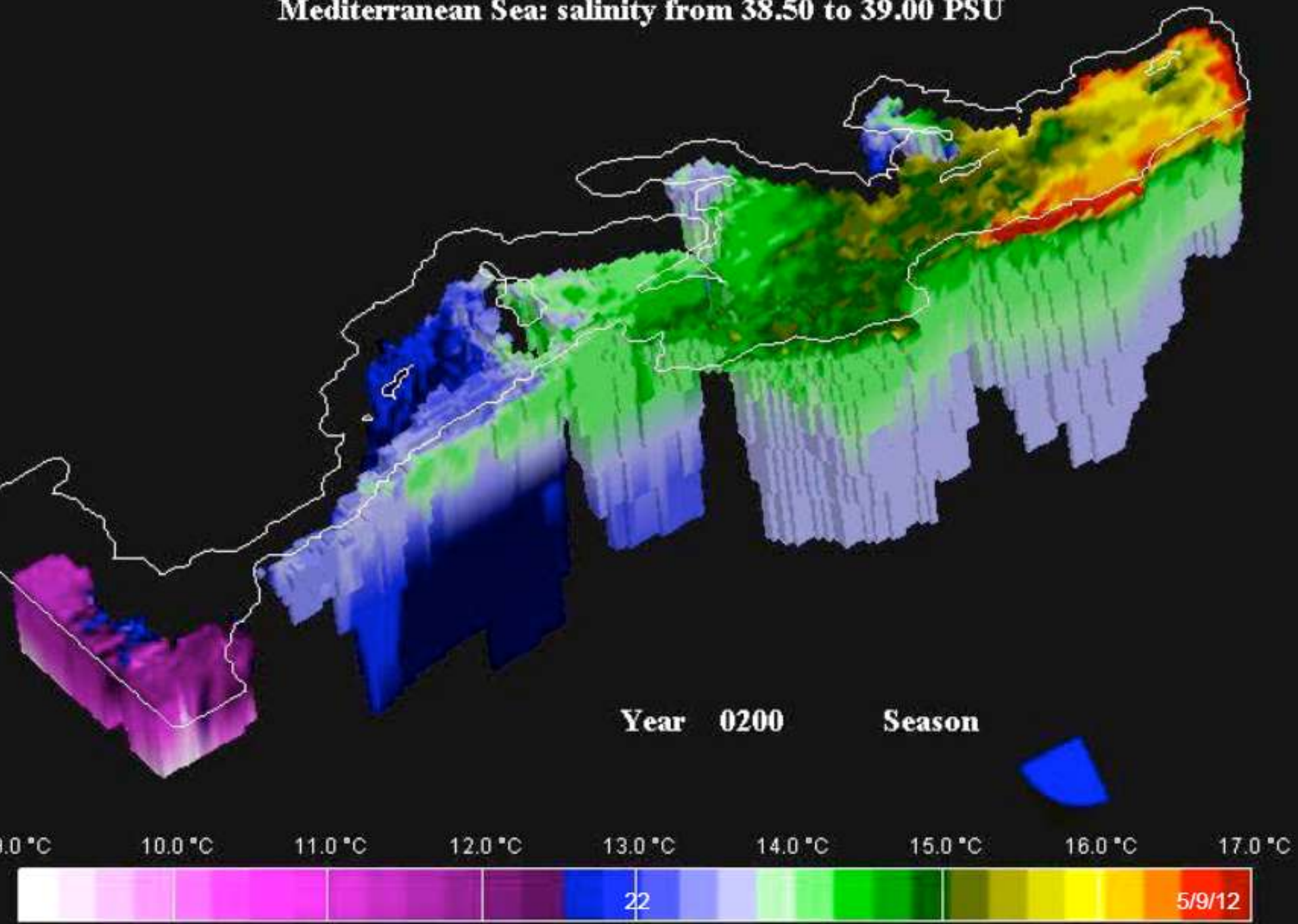
Evaporating column of seawater

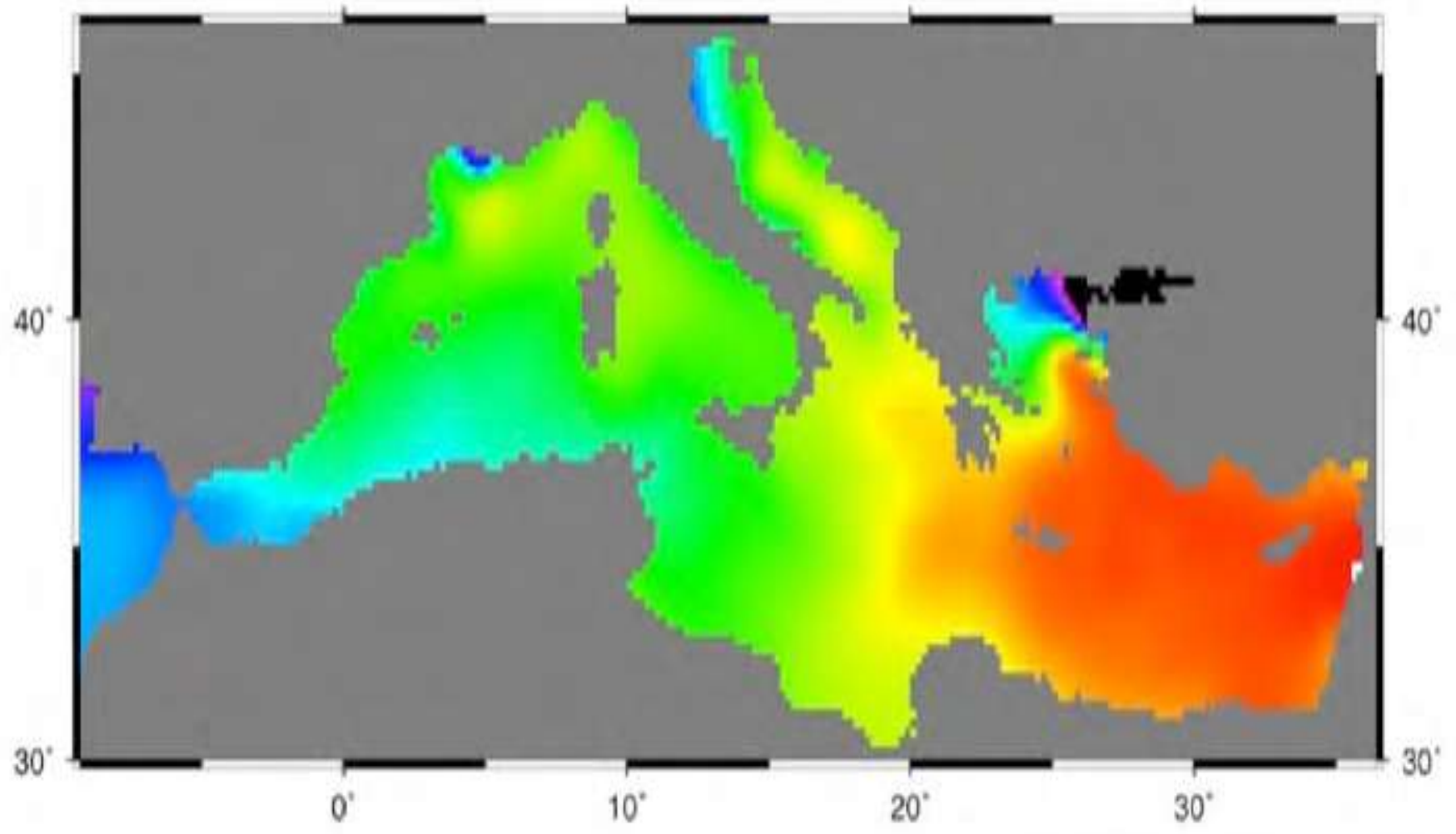
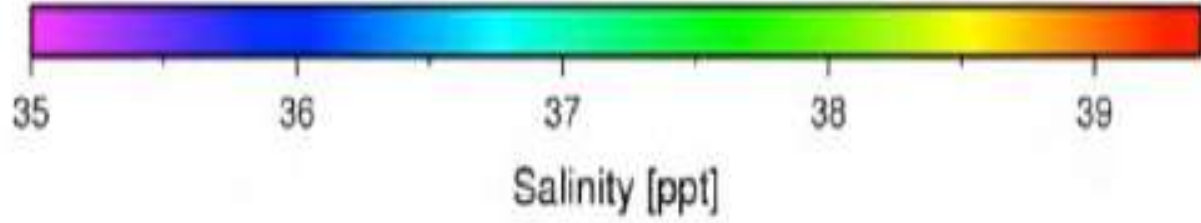


Luchttemperatuur boven de MZ



Atlantic Ocean: salinity from 35.50 to 36.50 PSU
Mediterranean Sea: salinity from 38.50 to 39.00 PSU

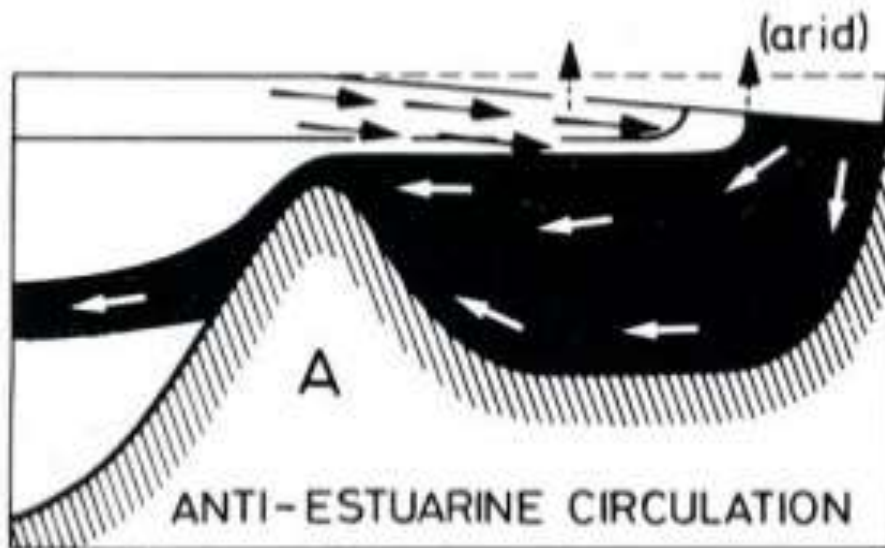




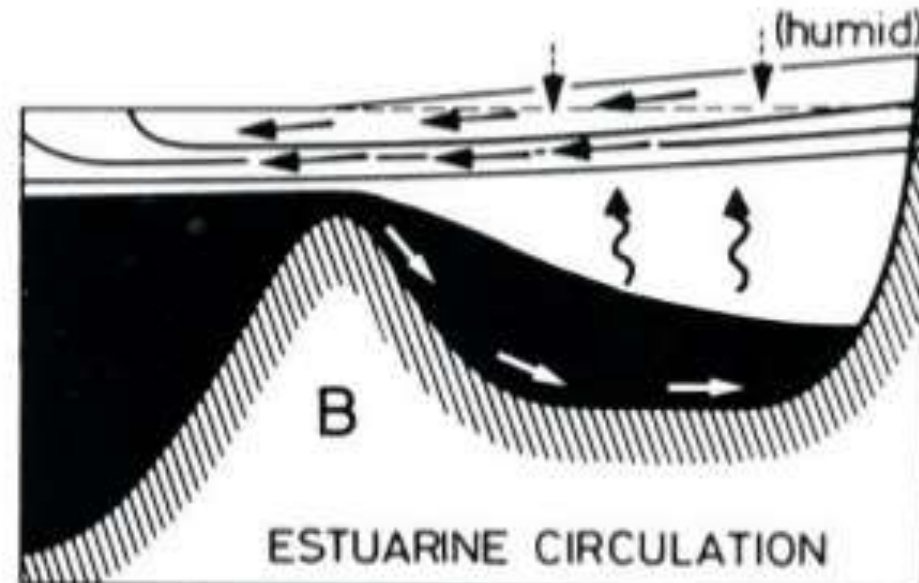
present-day sea surface salinity

Door klimaat een anti estuarine stroming in de MZ

Atlantic - Gibraltar - Mediterr.
Indic - Bab el Mandeb - Red Sea
Indic - Hormus - Pers. Arab. Gulf

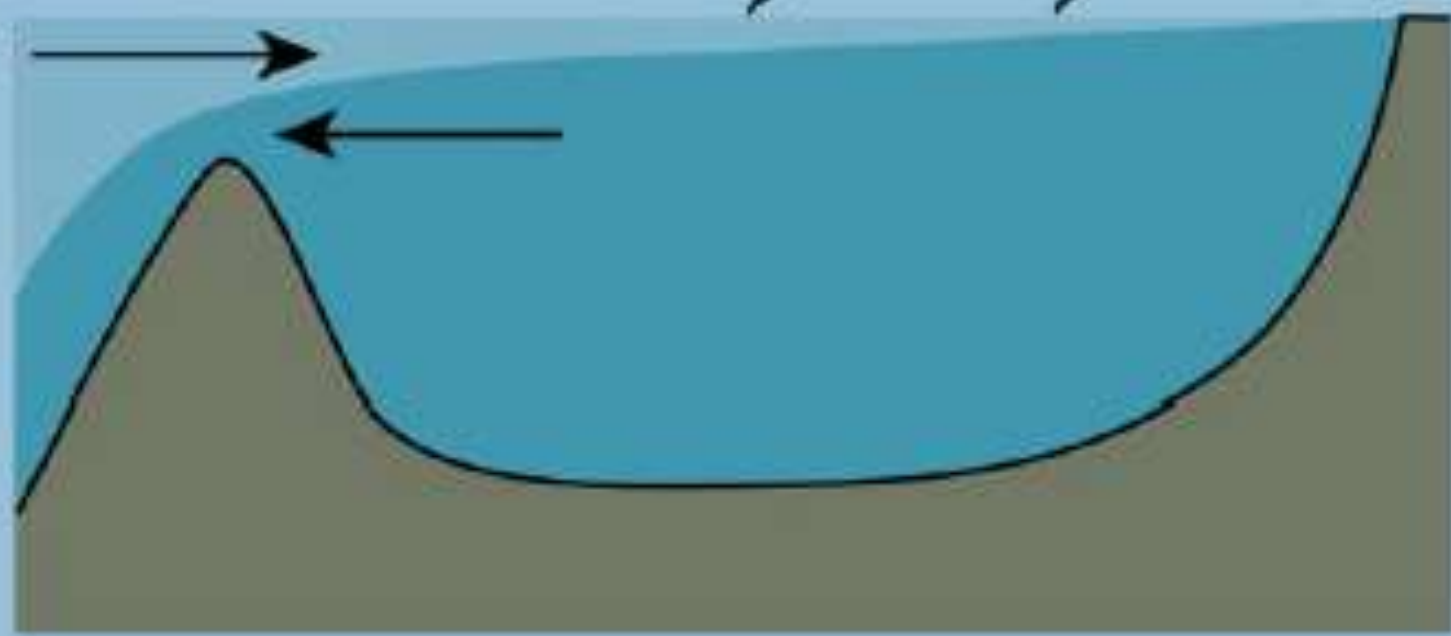


Atlantic-Sill-Norwegian and Greenland Fjords
Mediterr. - Bosporus - Black Sea
North Sea - Belts - Baltic



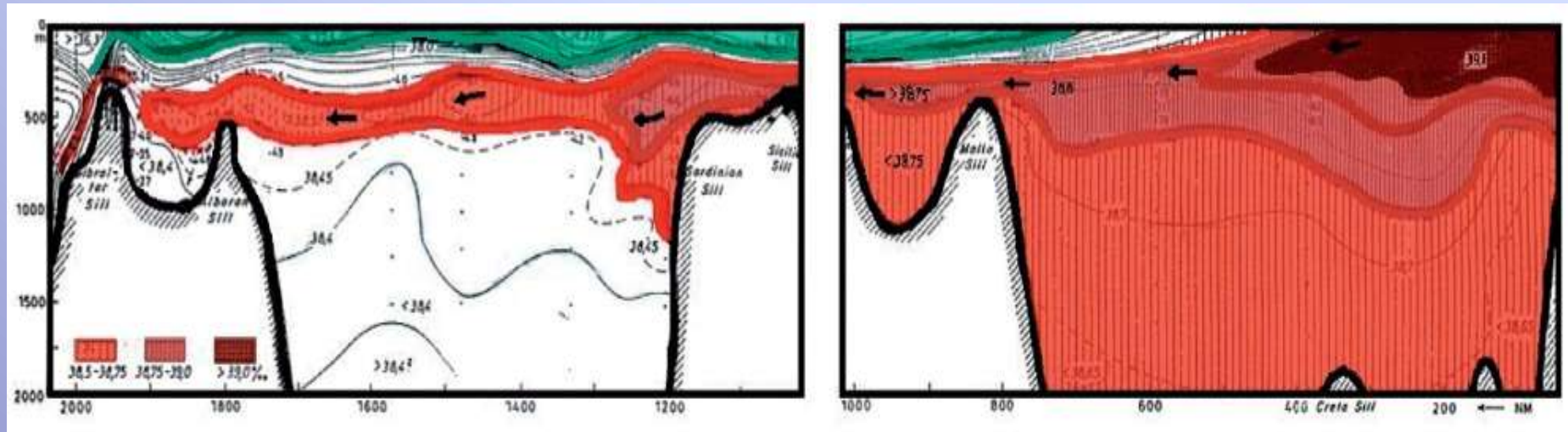
$$\approx 60.000 \text{ m}^3/\text{s} = 0.06 \text{ Sv}$$

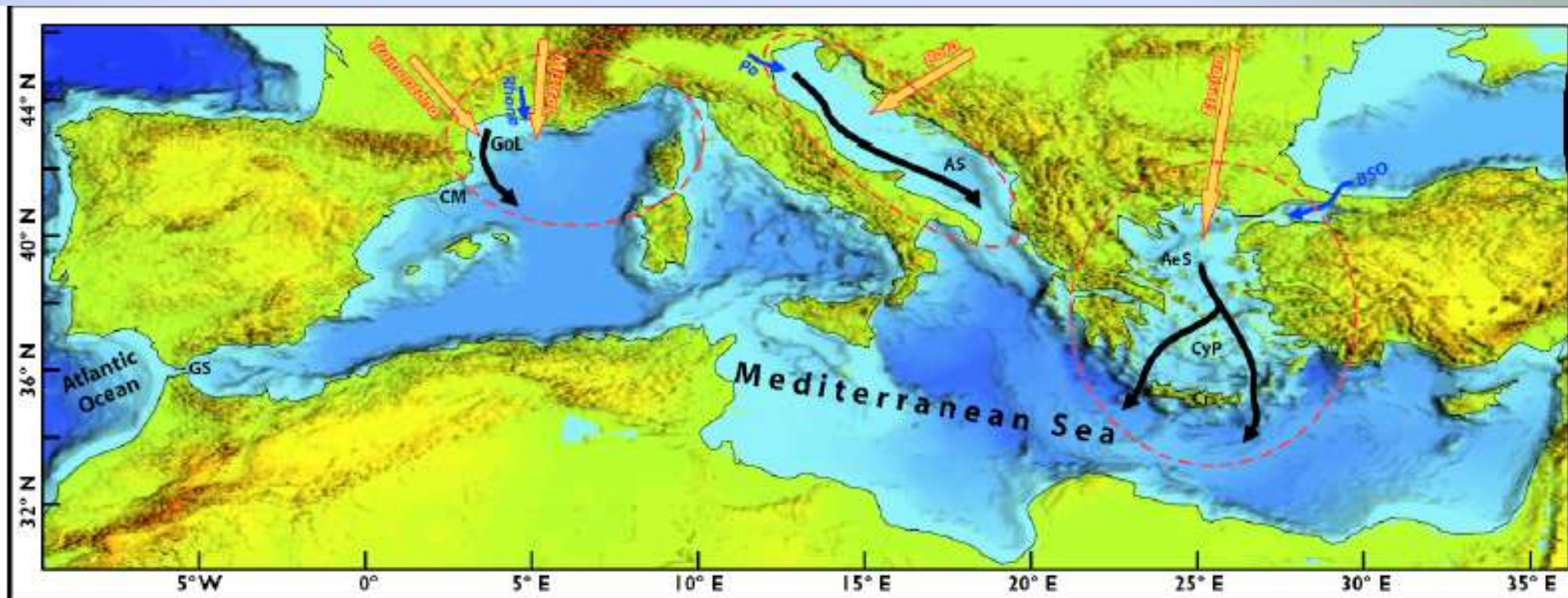
0.7-1.4 Sv



$$1 \text{ Sverdrup} = 10^6 \text{ m}^3/\text{s}$$

Zoutgehalte en stroming

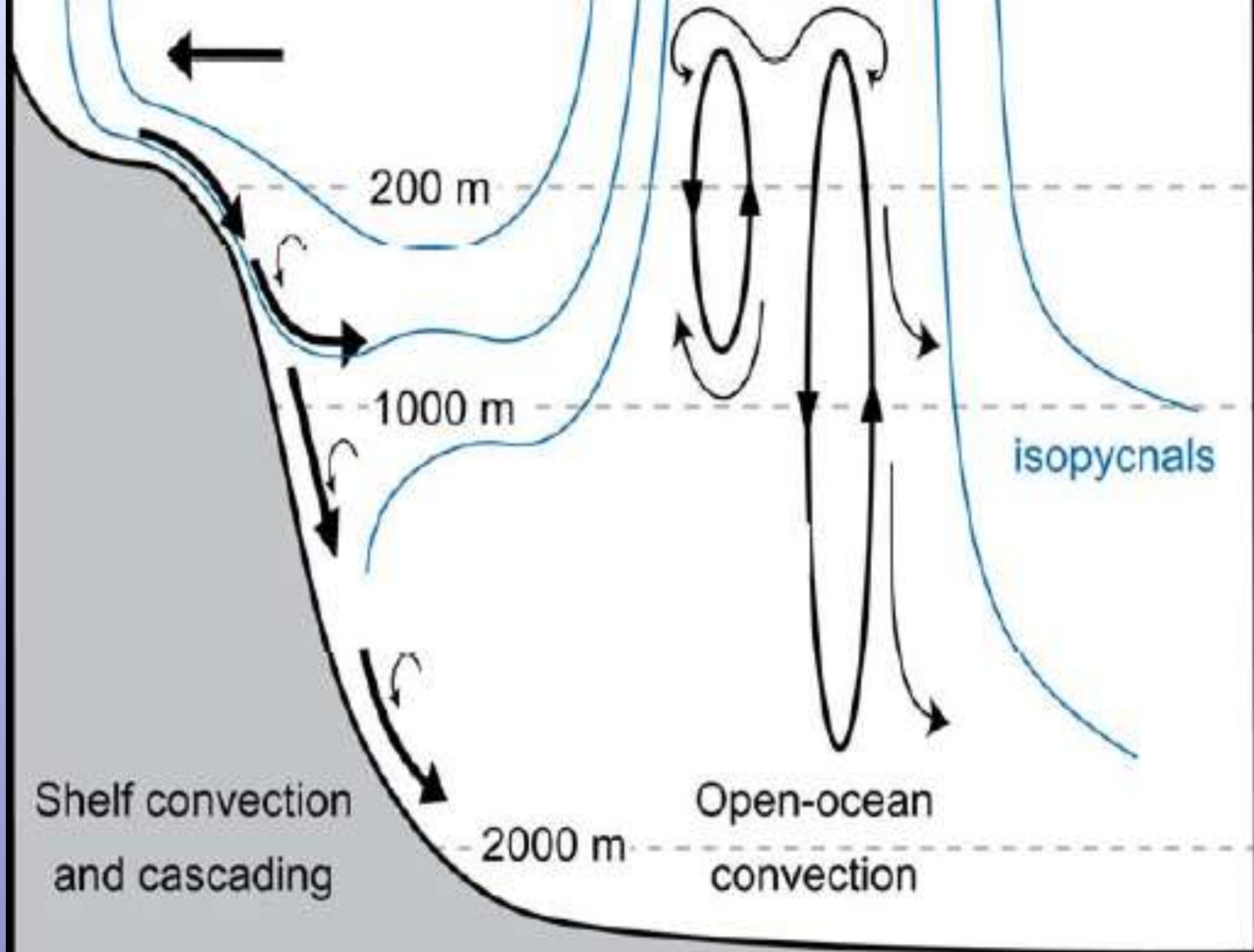


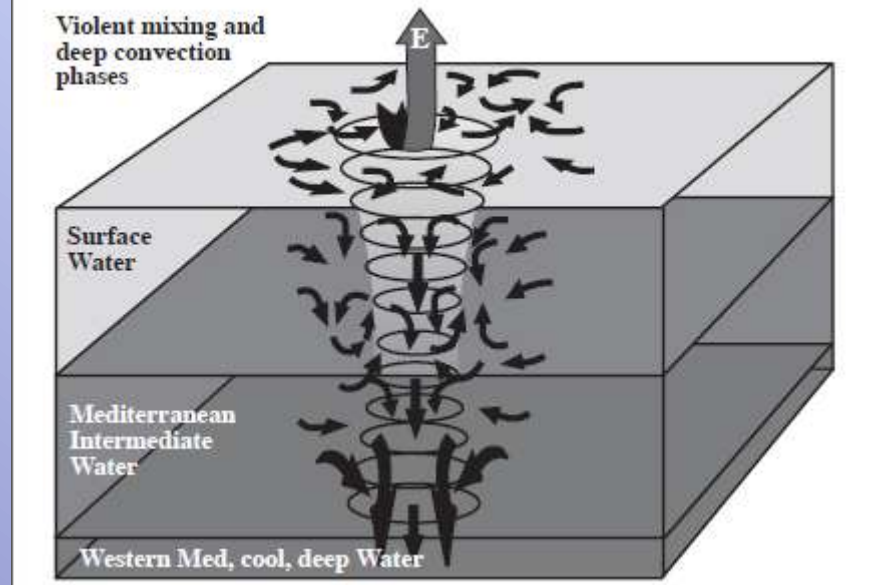
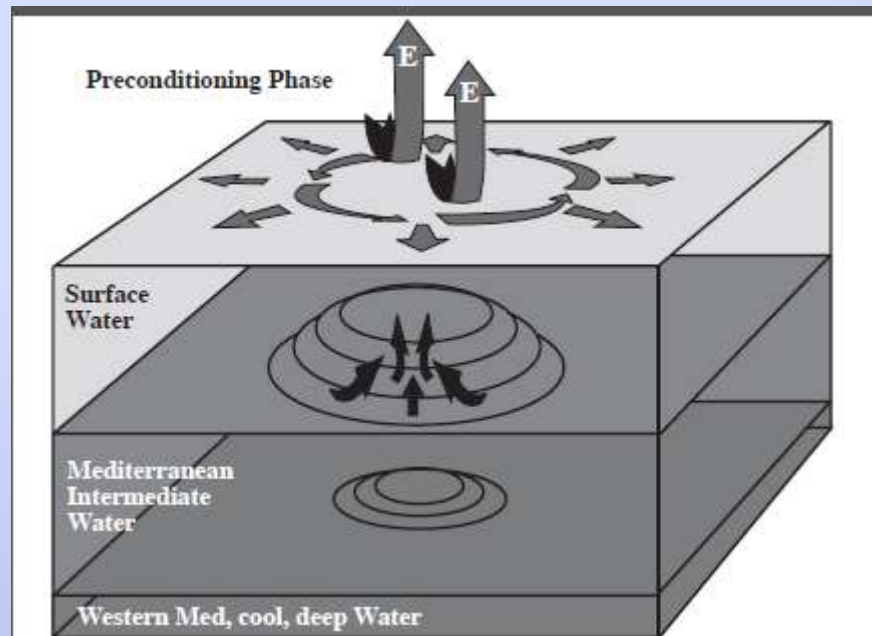


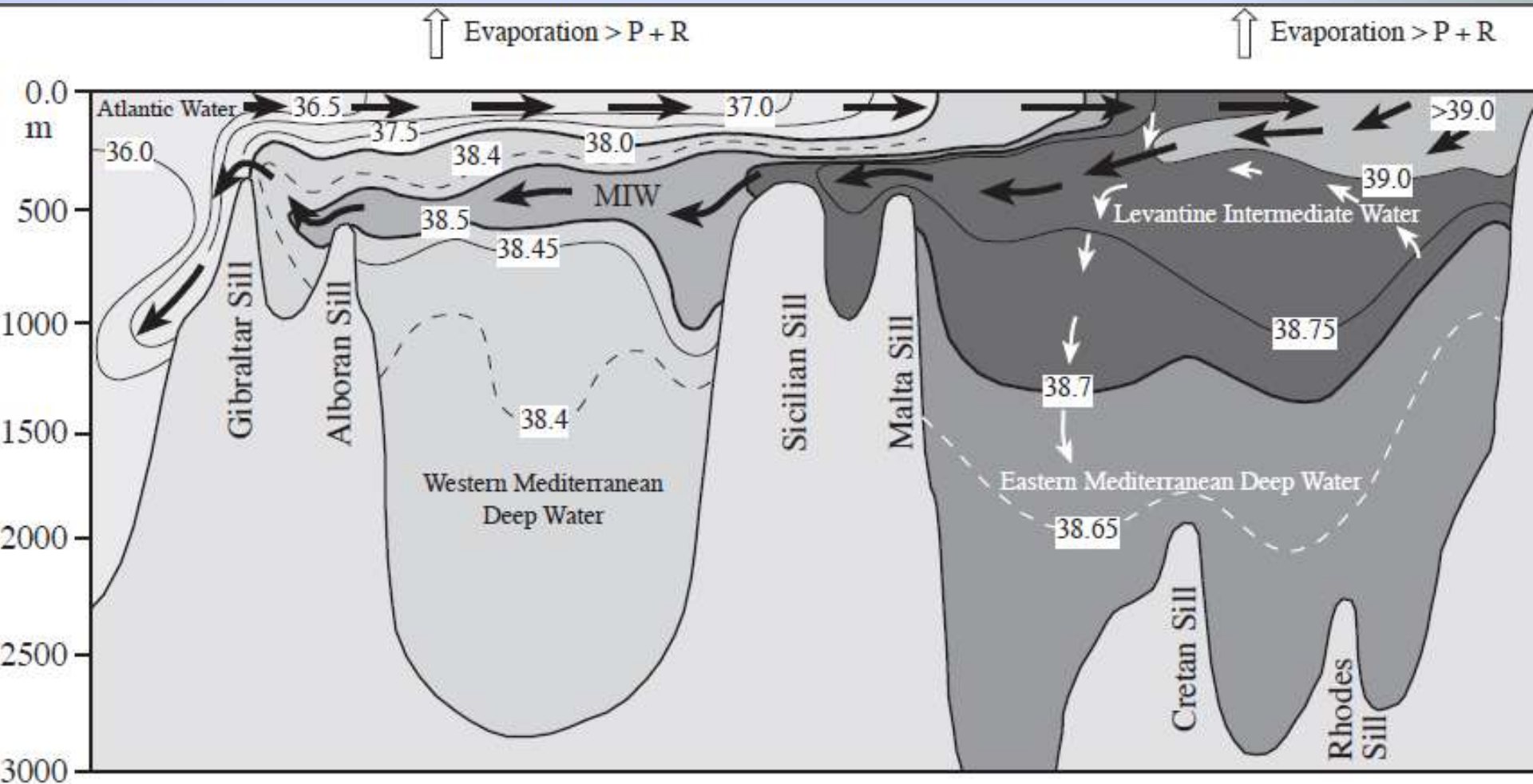
INVLOED RIVIEREN

windrichtingen

reliëf



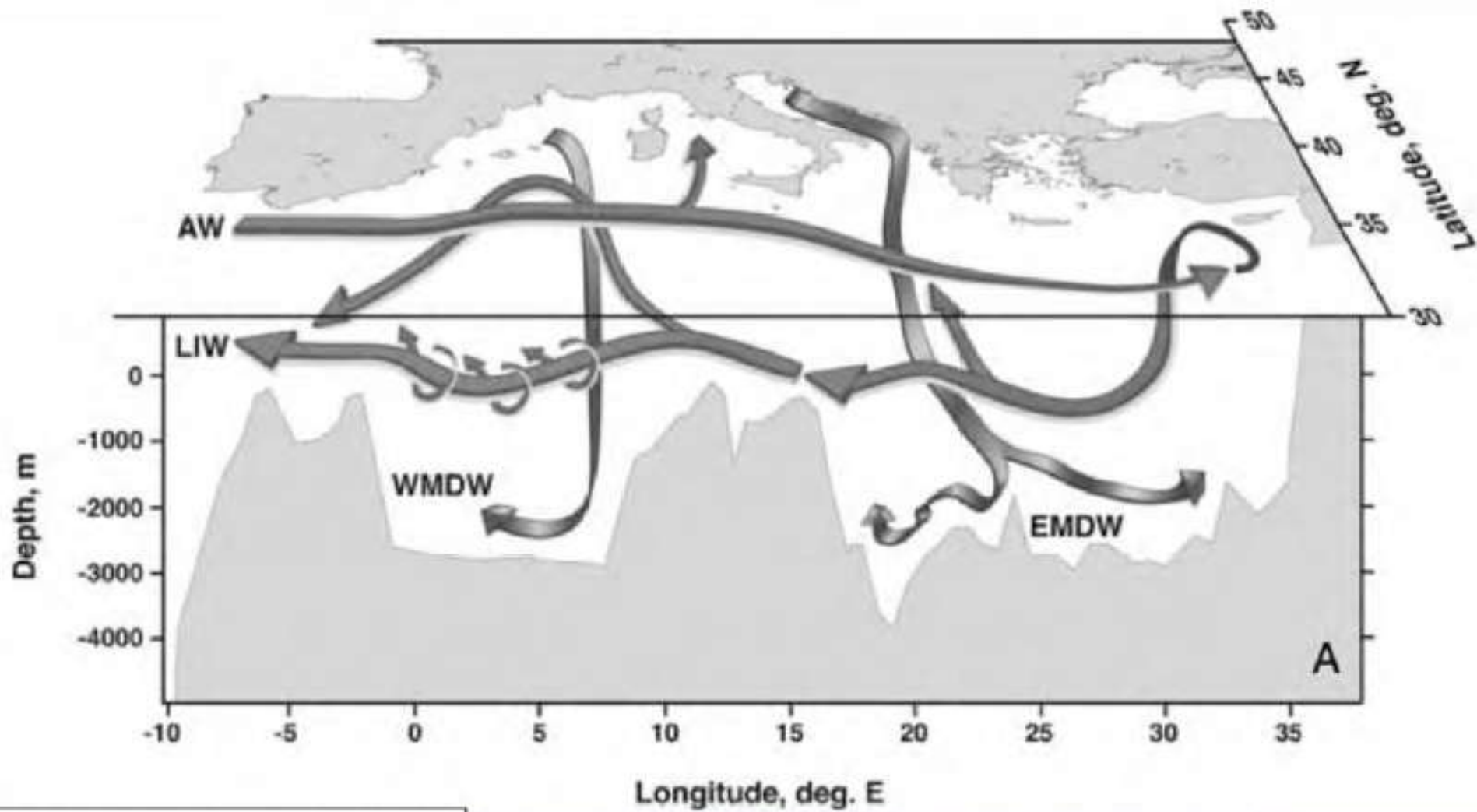




Enkele getallen/richtingen

- Verdamping boven MZ : 4.690 km³
- Aanvoer door rivieren en neerslag: 1.830 km³
- Tekort: 2.860 km³
- Aanvoer vanuit AO: 53.000 km³
- Afvoer naar AO: 50.000 km³
- Neerslag evaporieten in MZ : 10.000.000 km³

- Oppervlakte MZ: 2,5 milj. km²

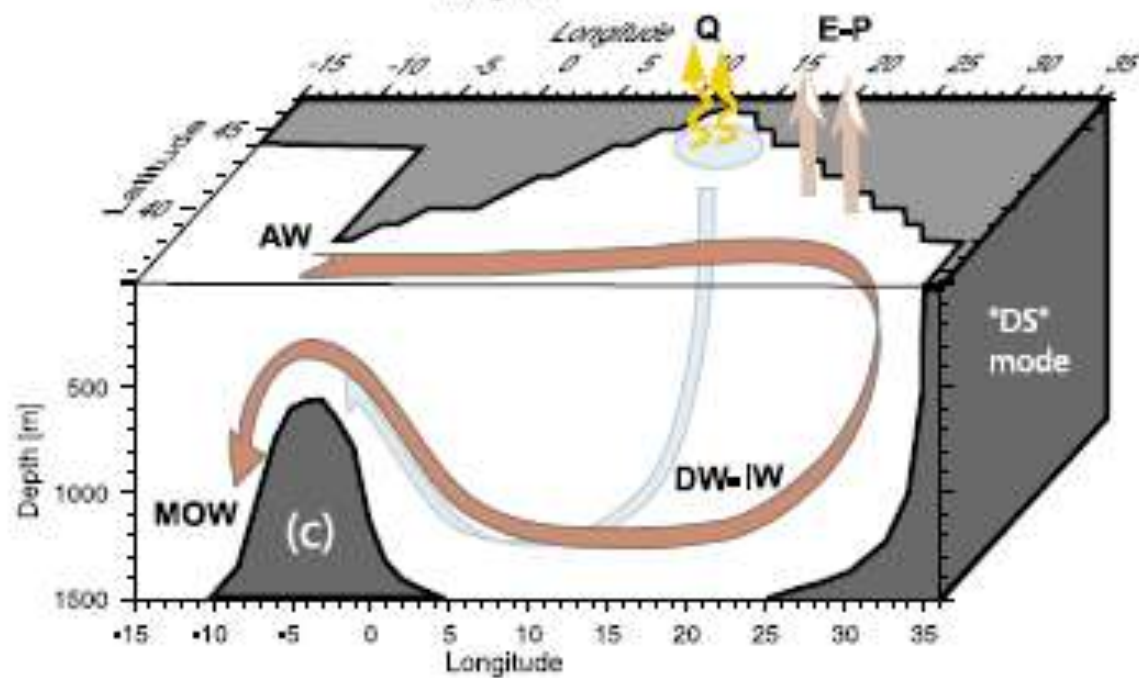
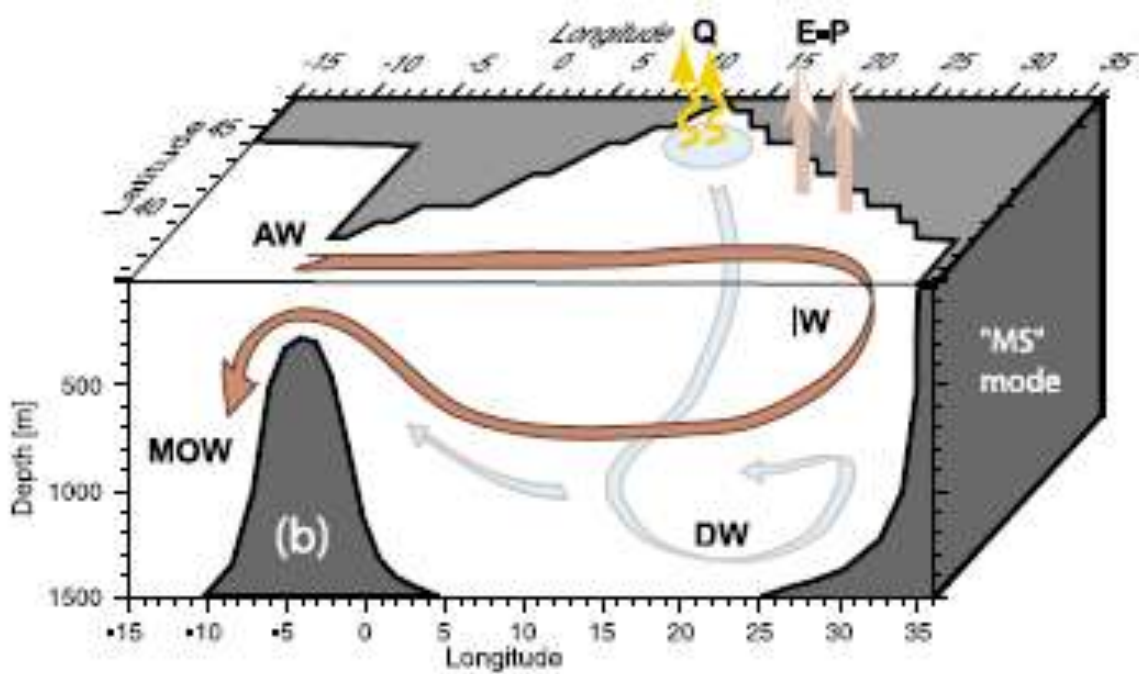


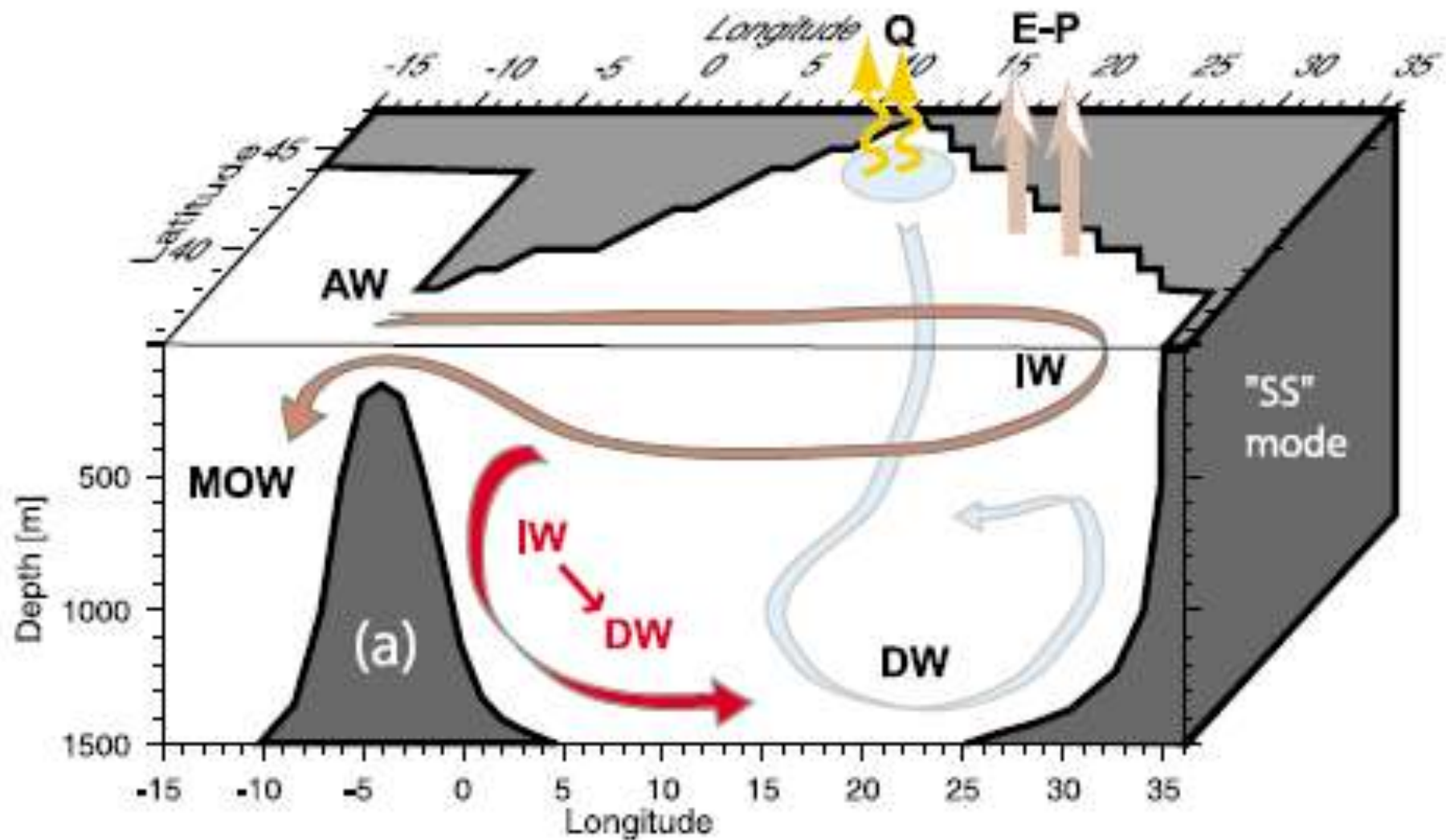
thermohaliene circulatie

Tsimplis et al. (2005)

Onderzoek naar invloed drempelhoogte op stromingen

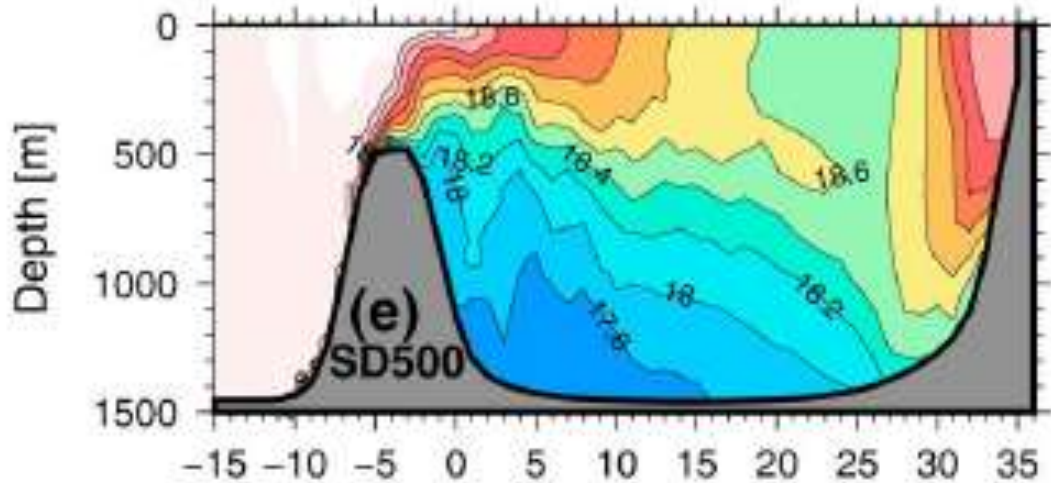
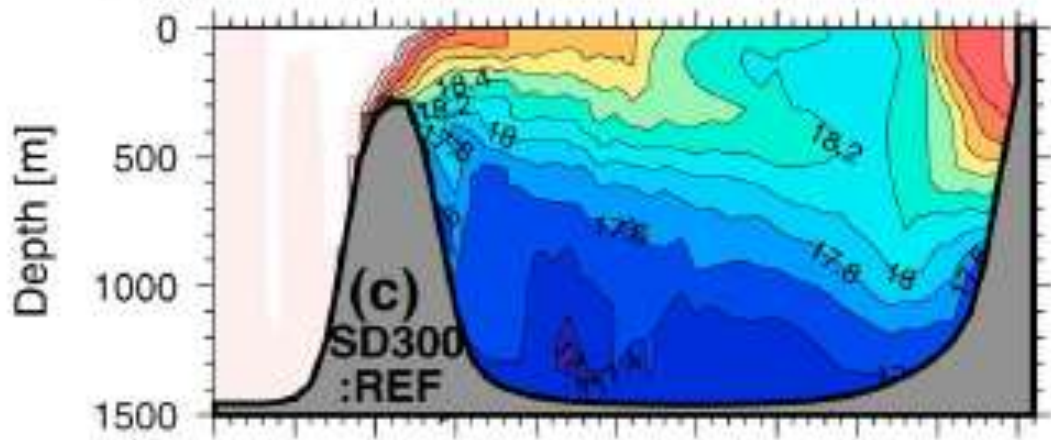
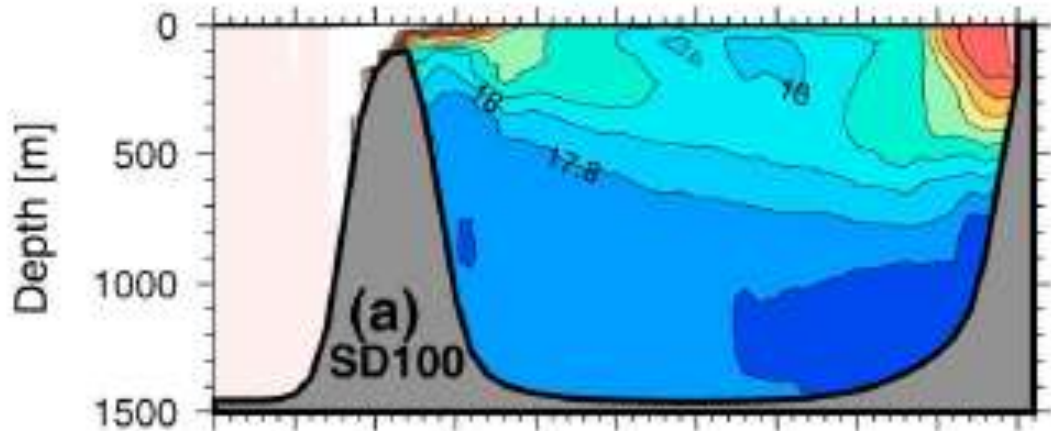
- Moderate sill (zoals nu)
- Deep sill (ver voor het Messinien)
- Shallow sill (tijdens zoutcrisis)

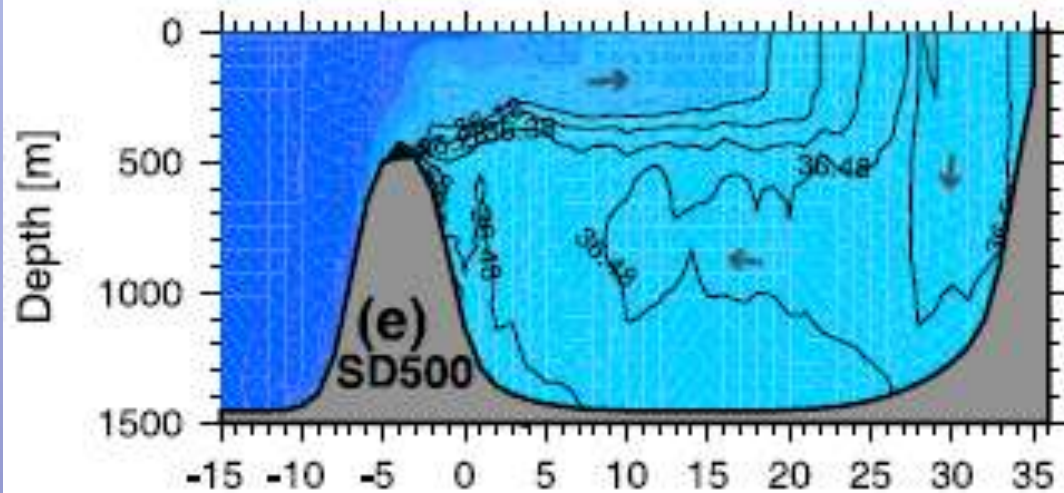
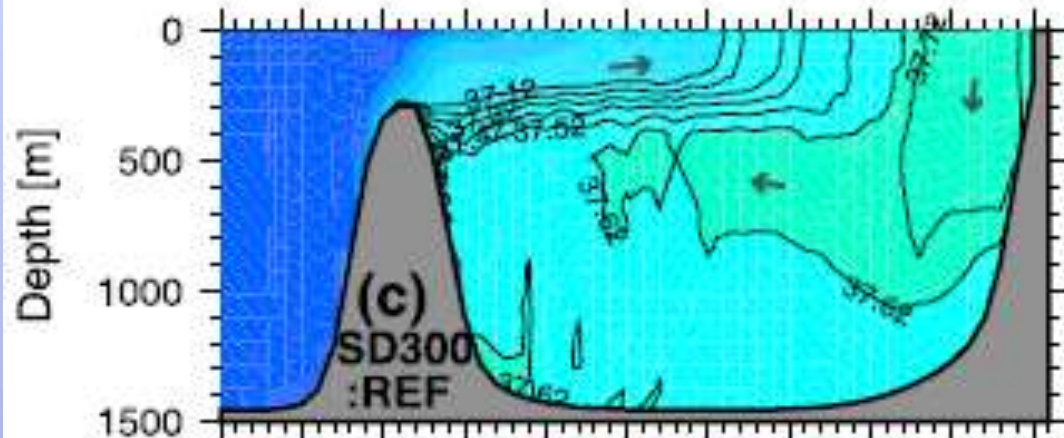
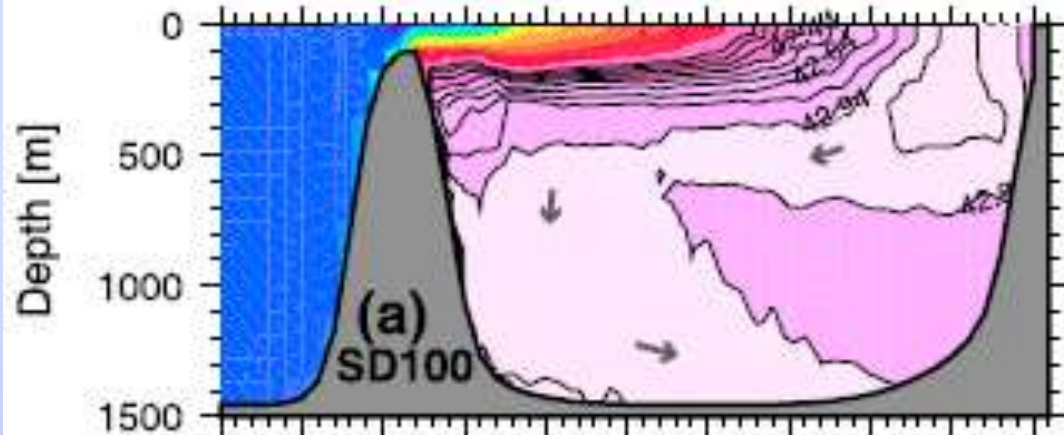


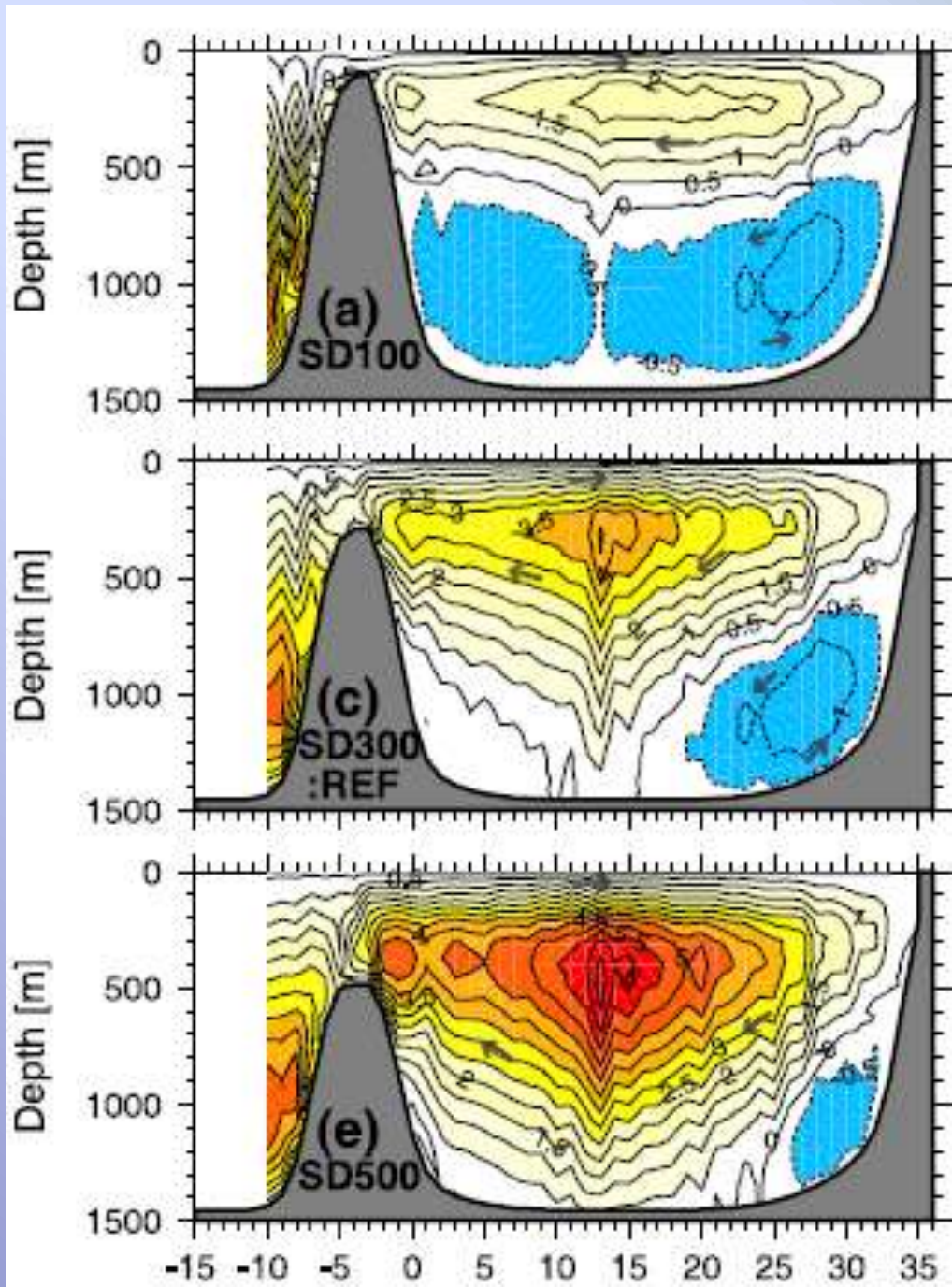


Invloed afname drempelhoogte (shoaling)

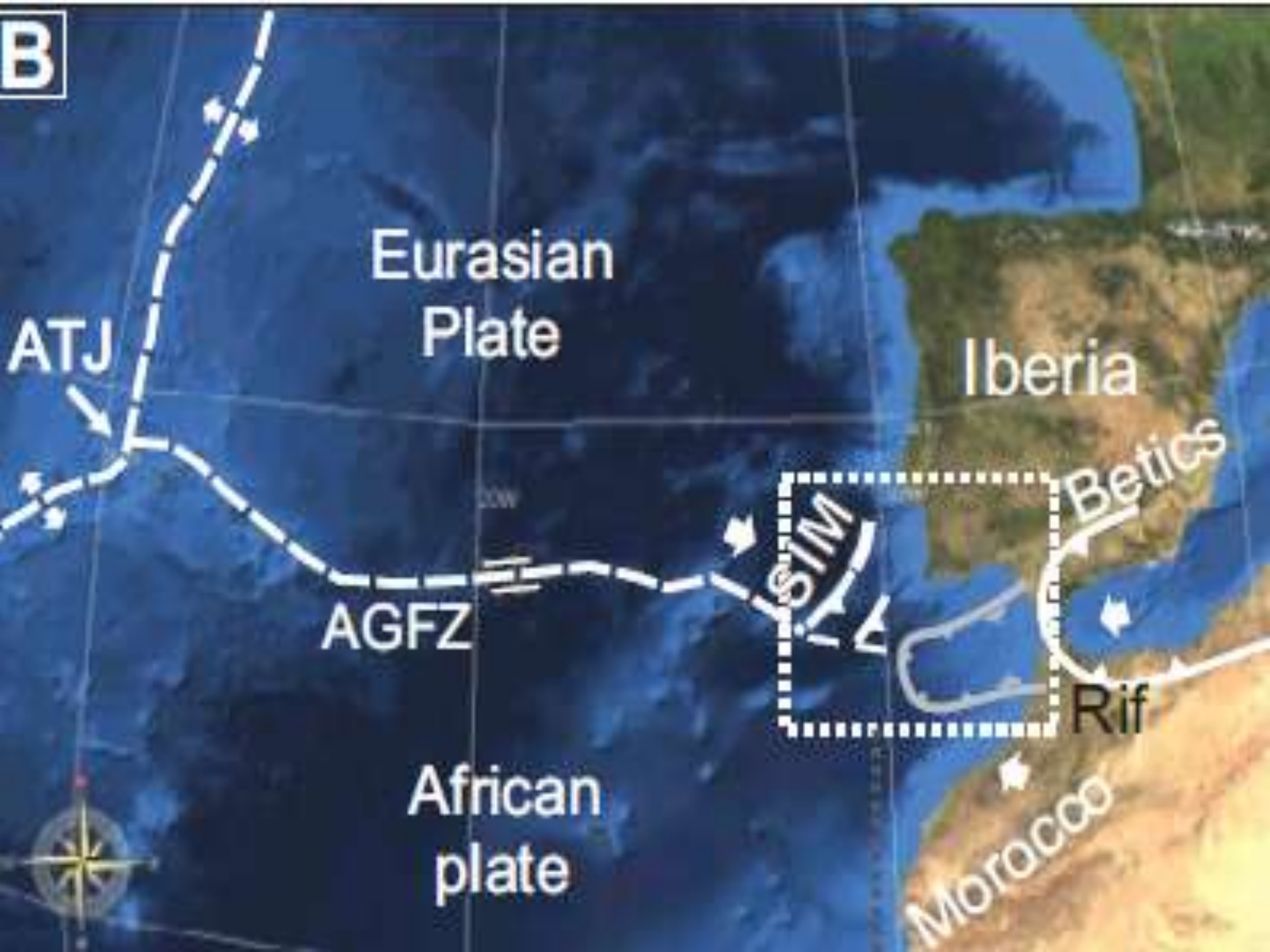
- De sterkte van de bovenstromingen verminderen
- Uitstroom raakt geblokkeerd hierdoor sterkere circulatie in onderste lagen
- Menging van water wordt minder, vooral in de diepere lagen

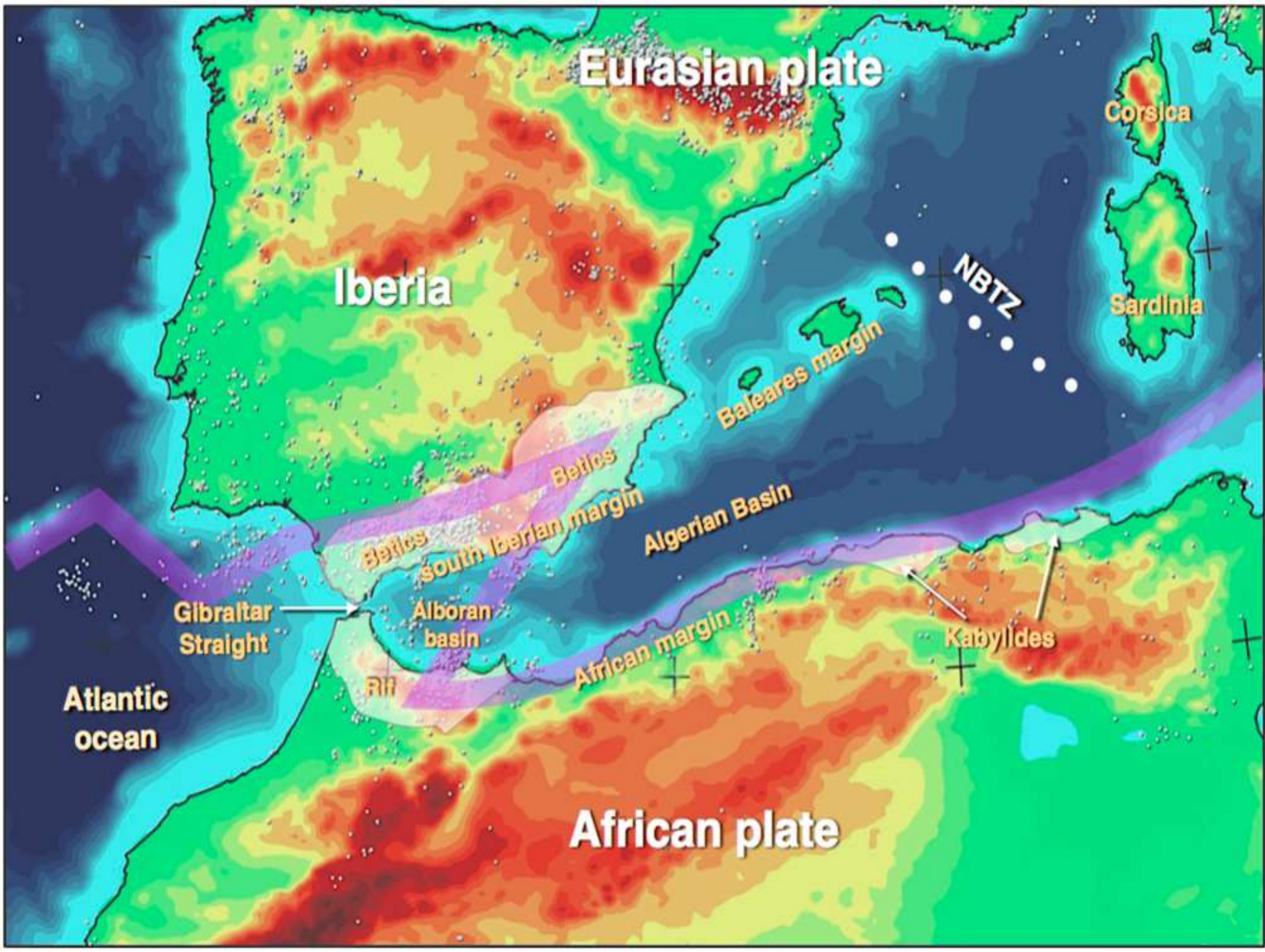








B



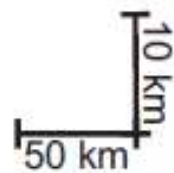
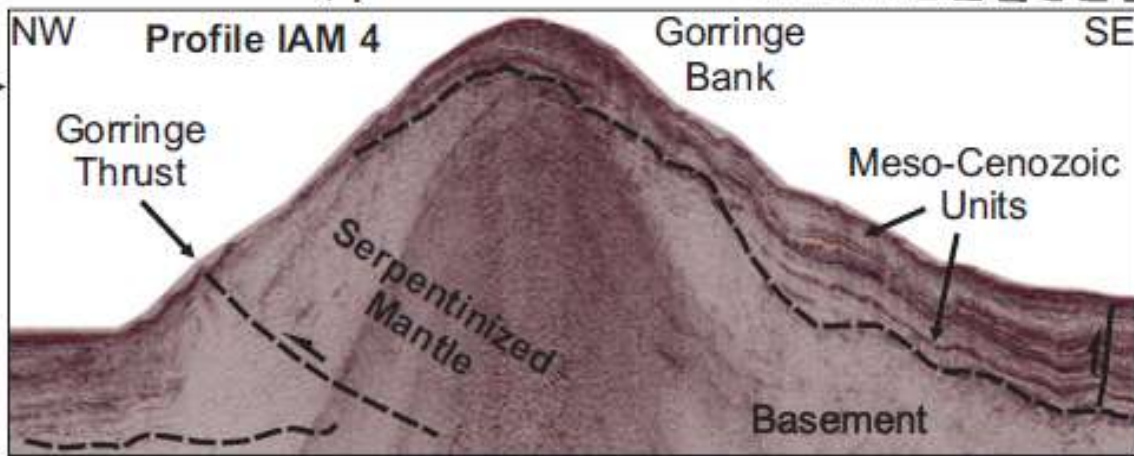
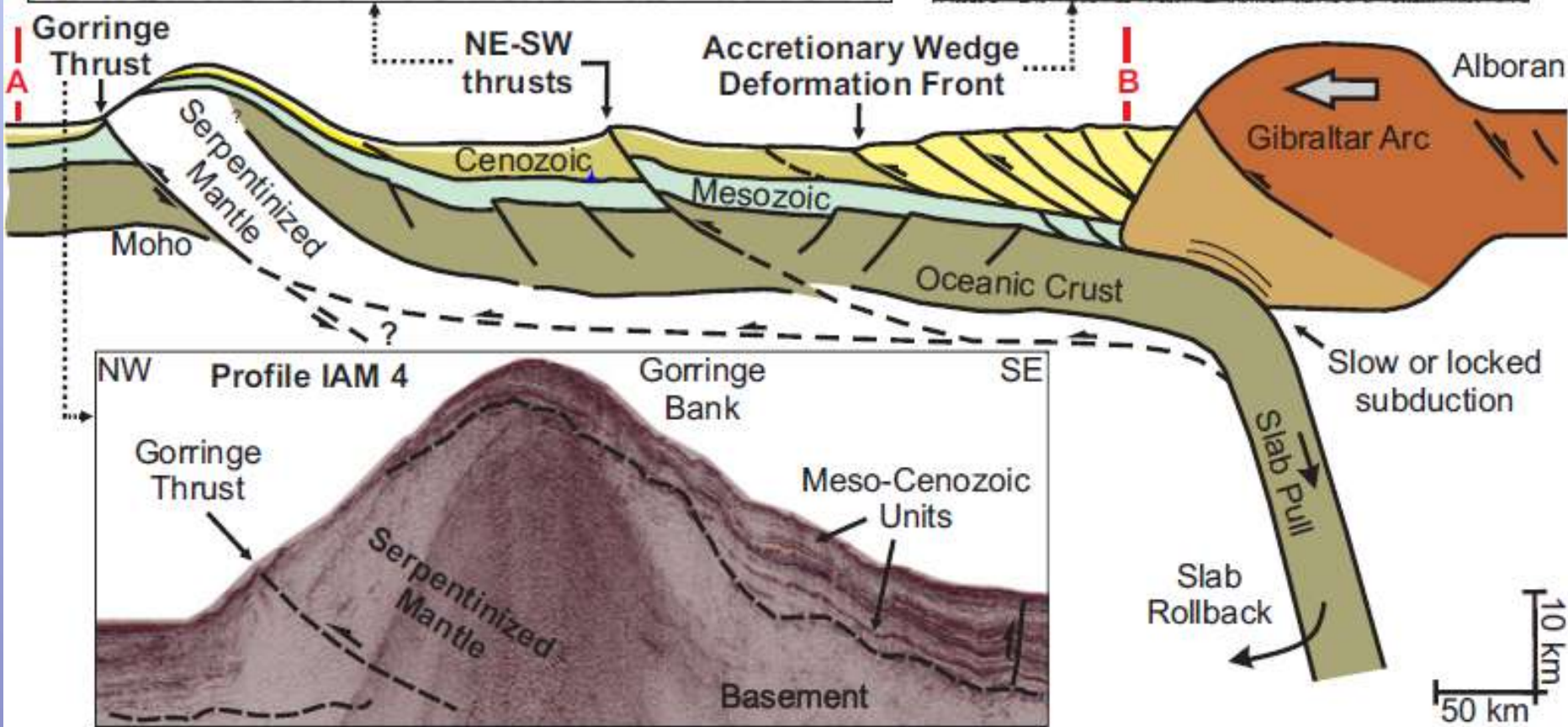
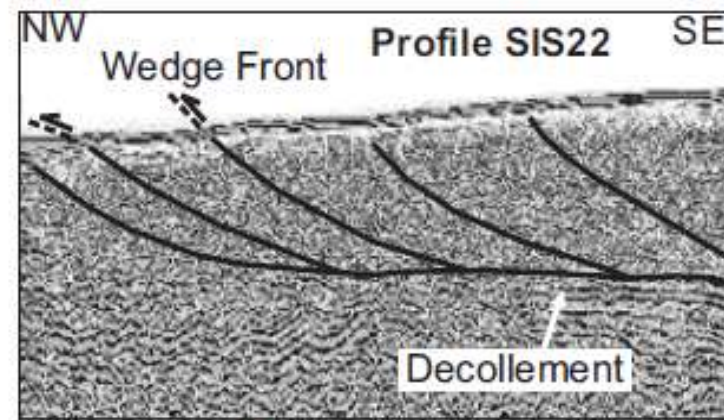
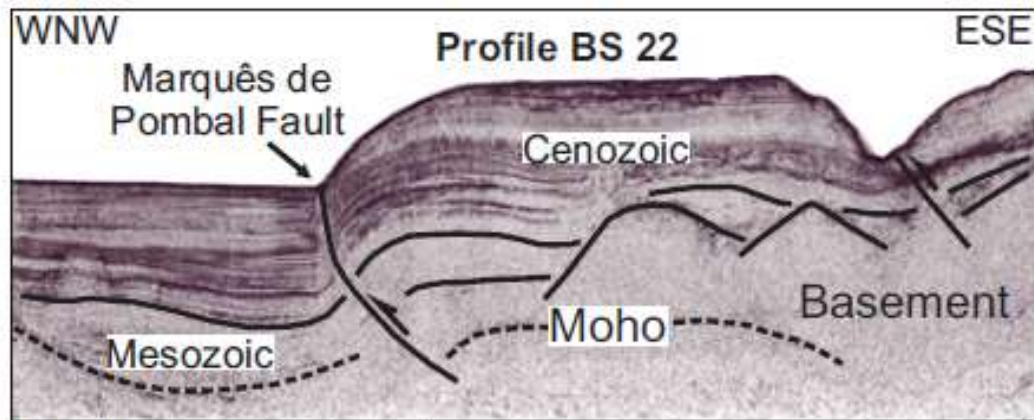
Welke processen spelen een rol in de Alboran regio?

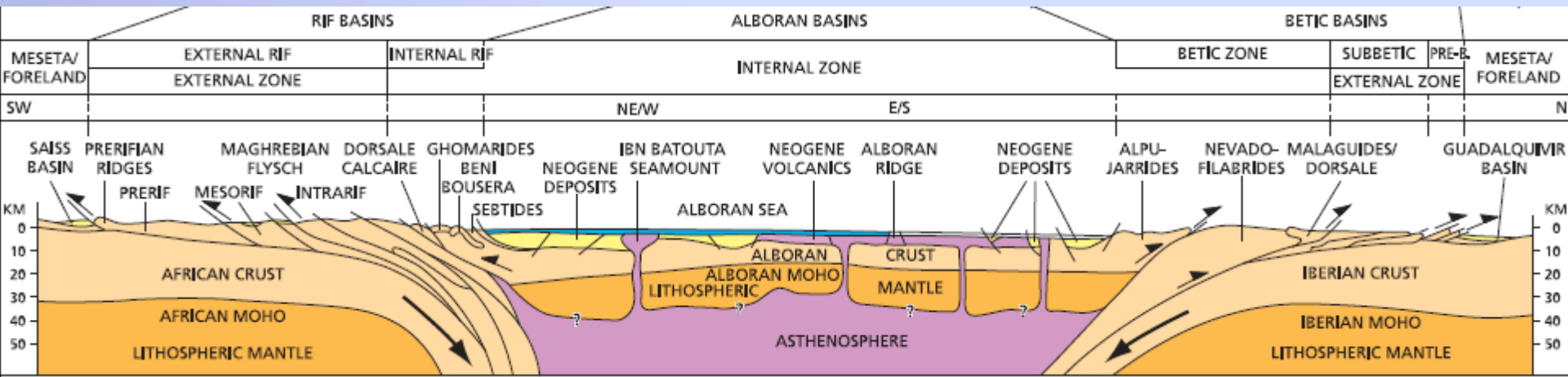
Convergentie

Subductie

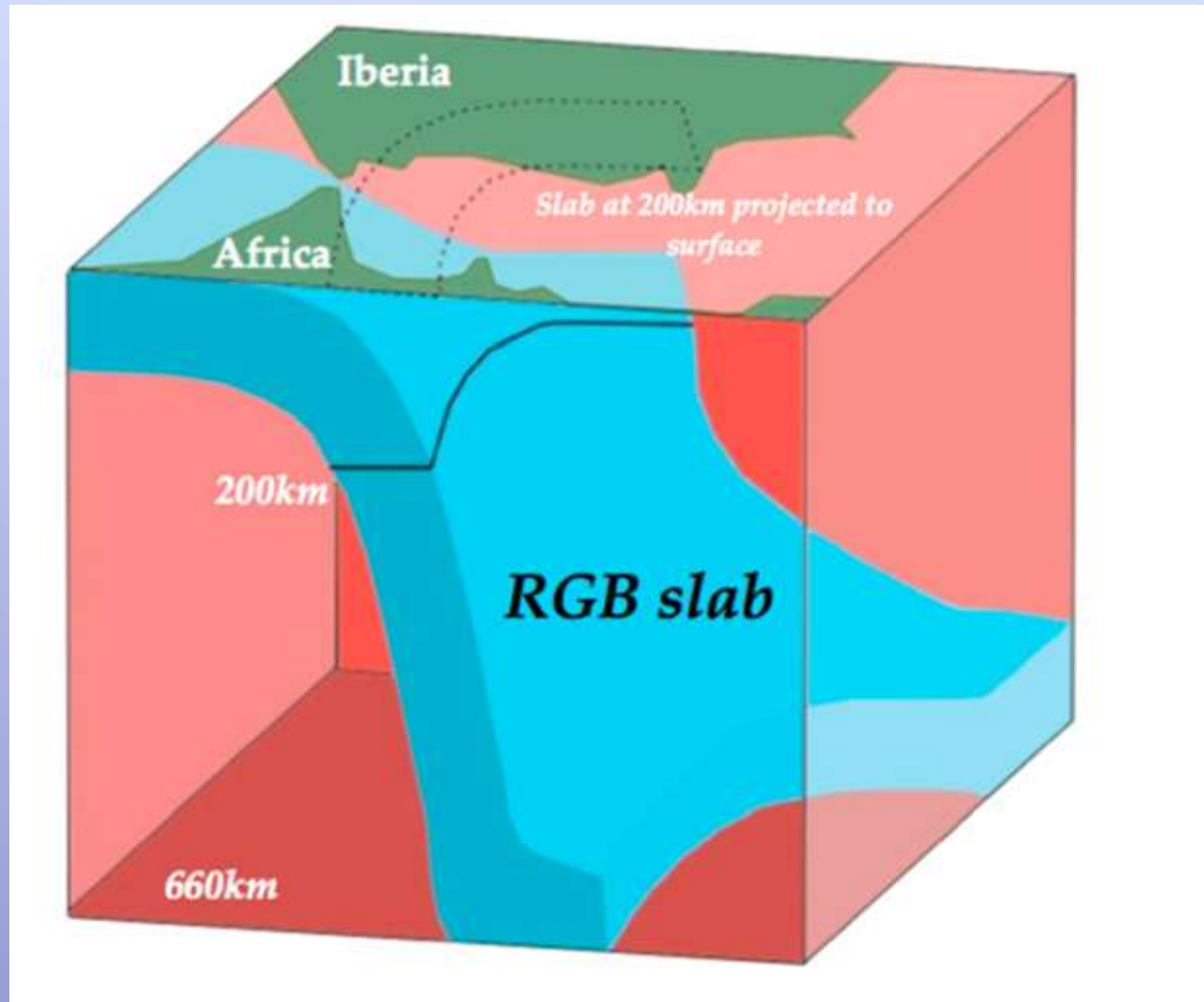
Slab roll back

Delaminatie?





De Rif Gibraltar Betic Slab

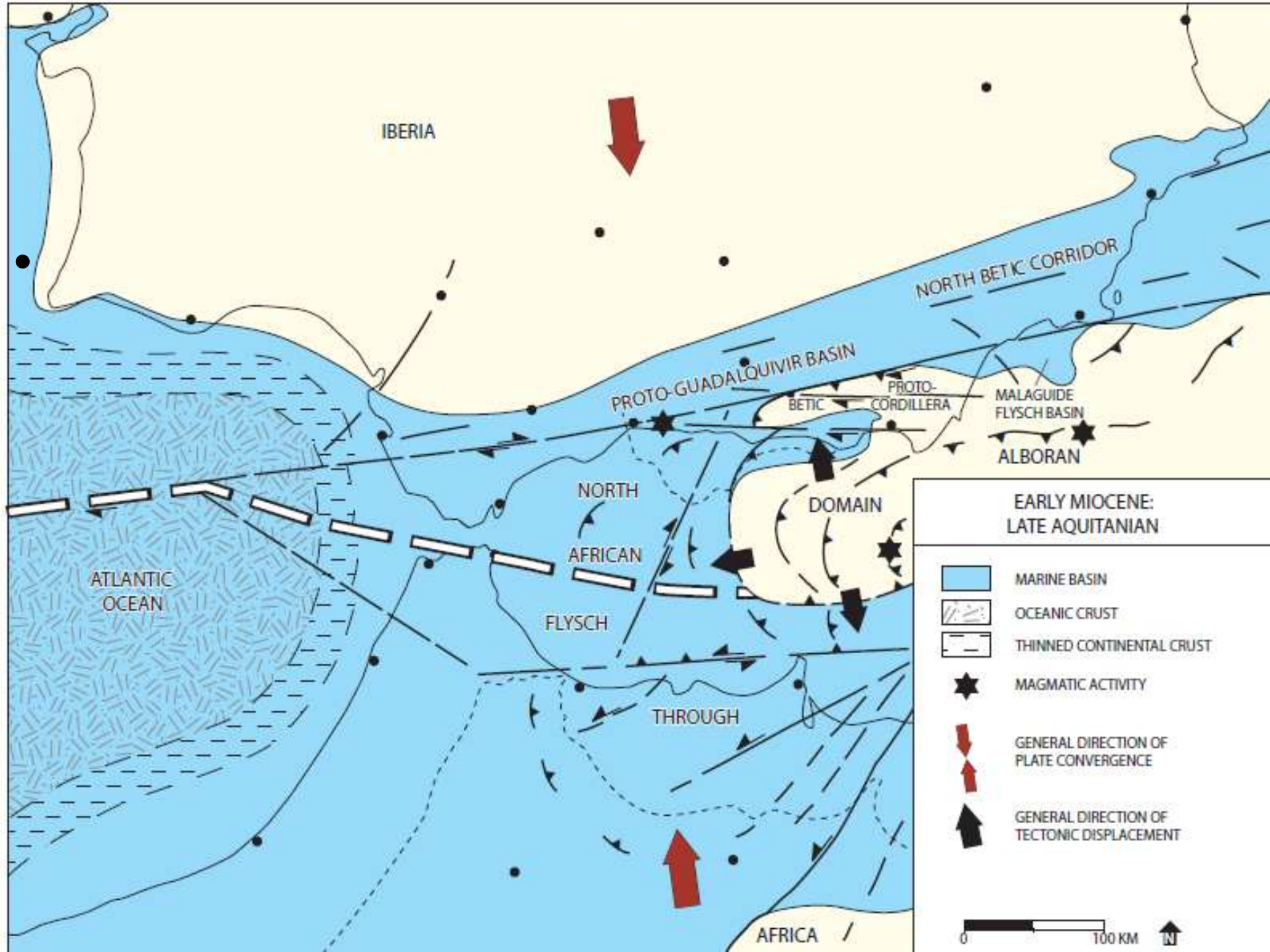


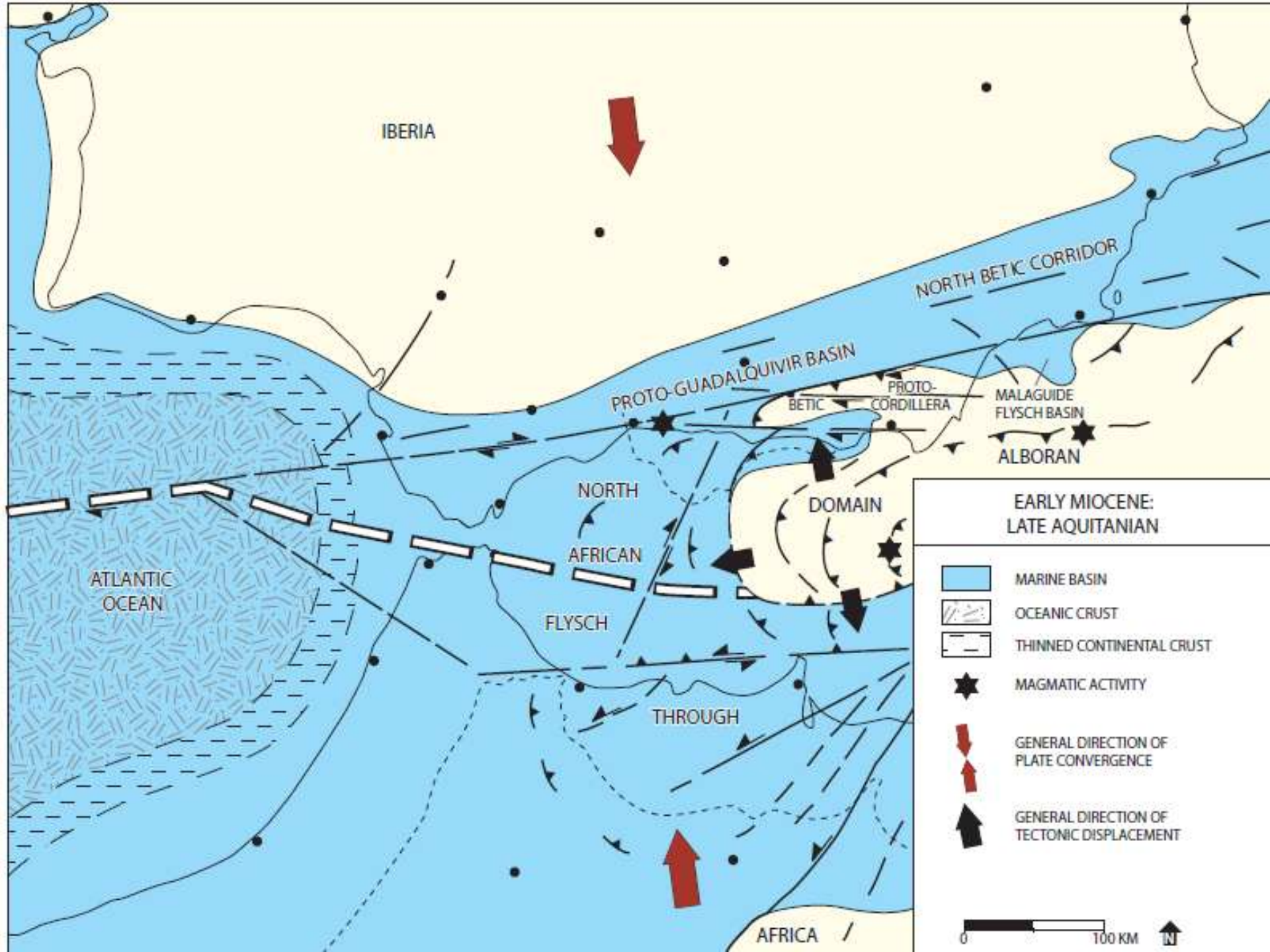
Fantasie?

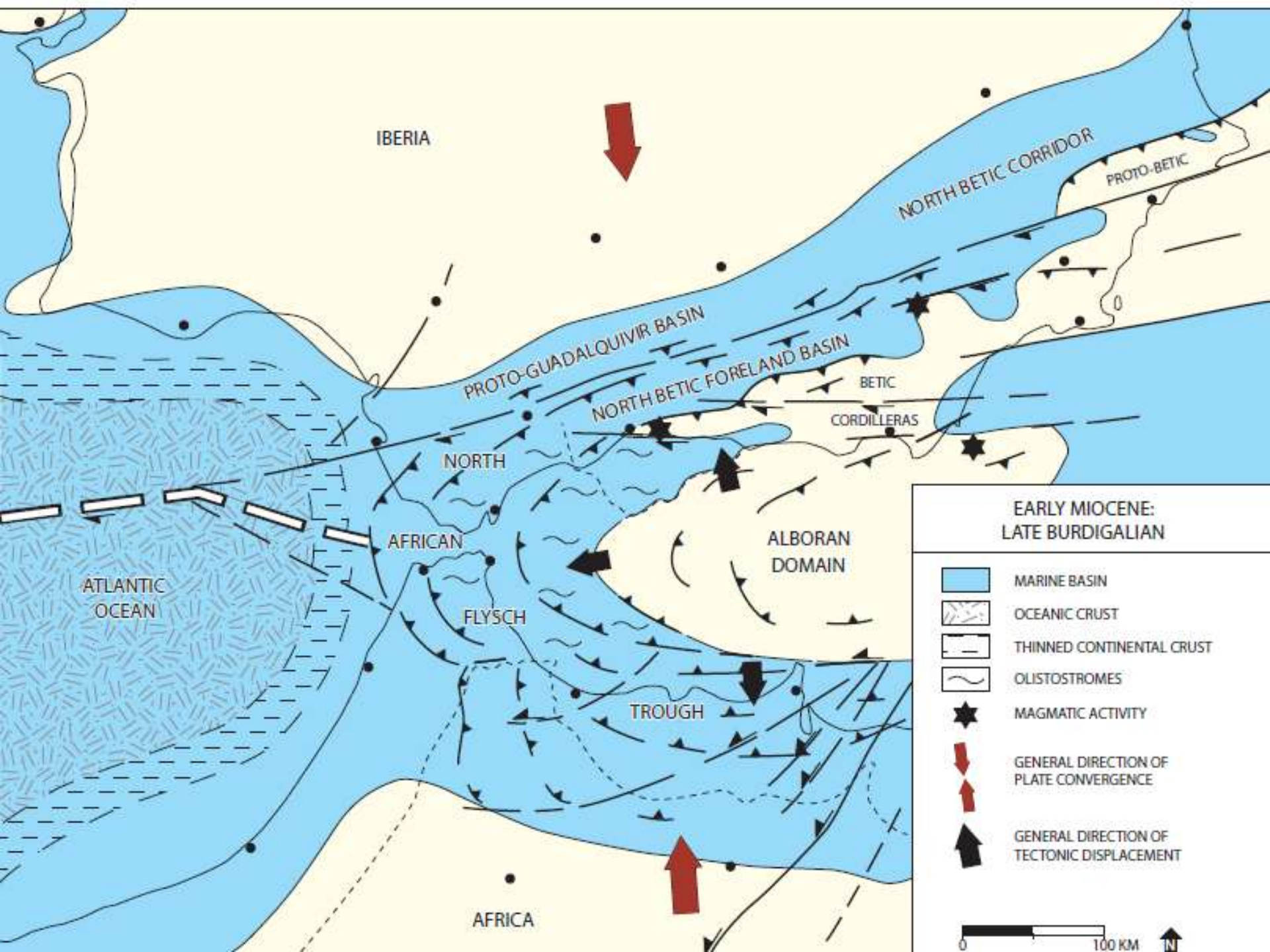


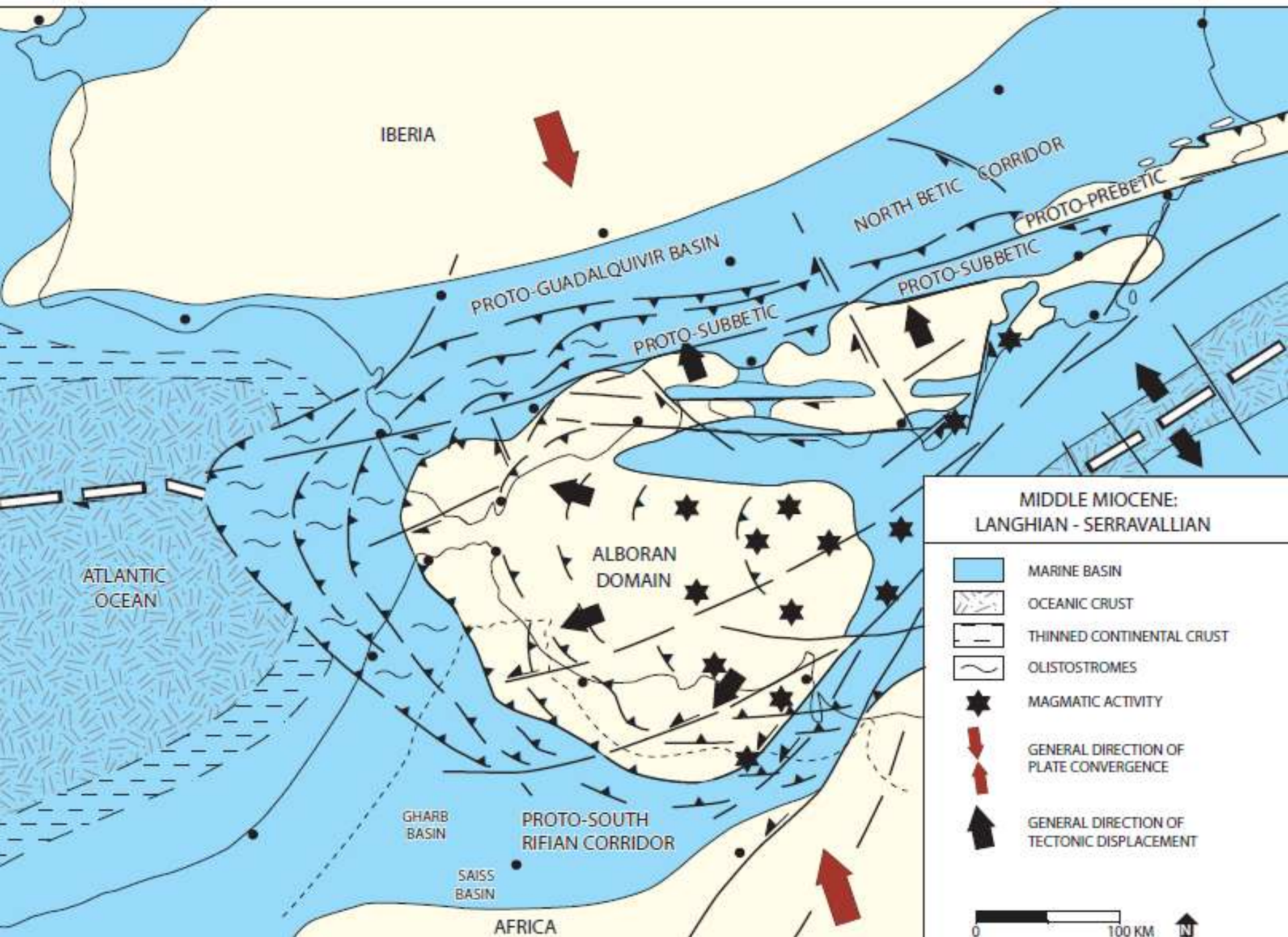
De zeestraten in zuidelijk Spanje en noordelijk Marokko











IBERIA

PROTO-GUADALQUIVIR BASIN

NORTH BETIC CORRIDOR

PROTO-PREBETIC

PROTO-SUBBETIC

ATLANTIC OCEAN

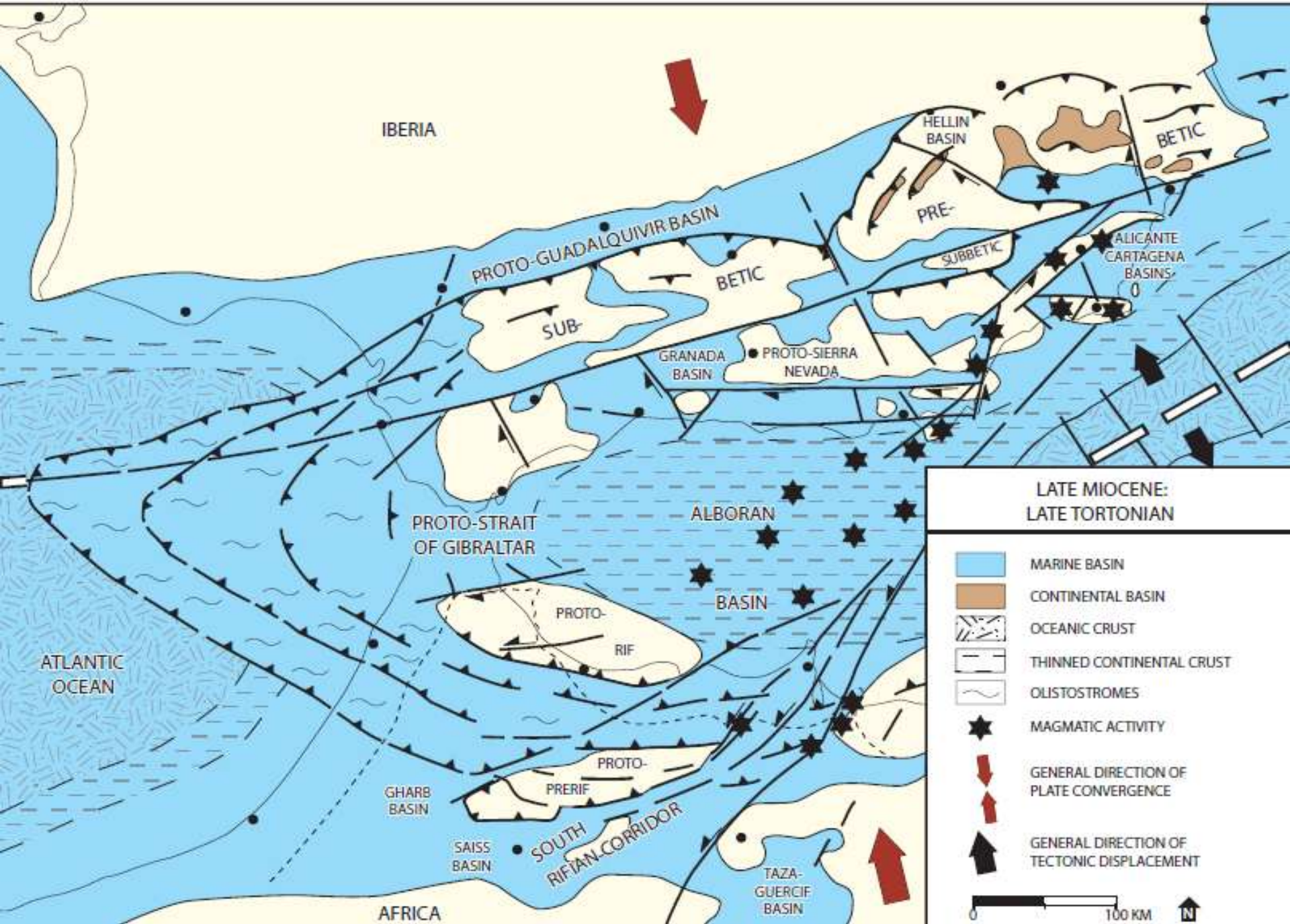
ALBORAN DOMAIN

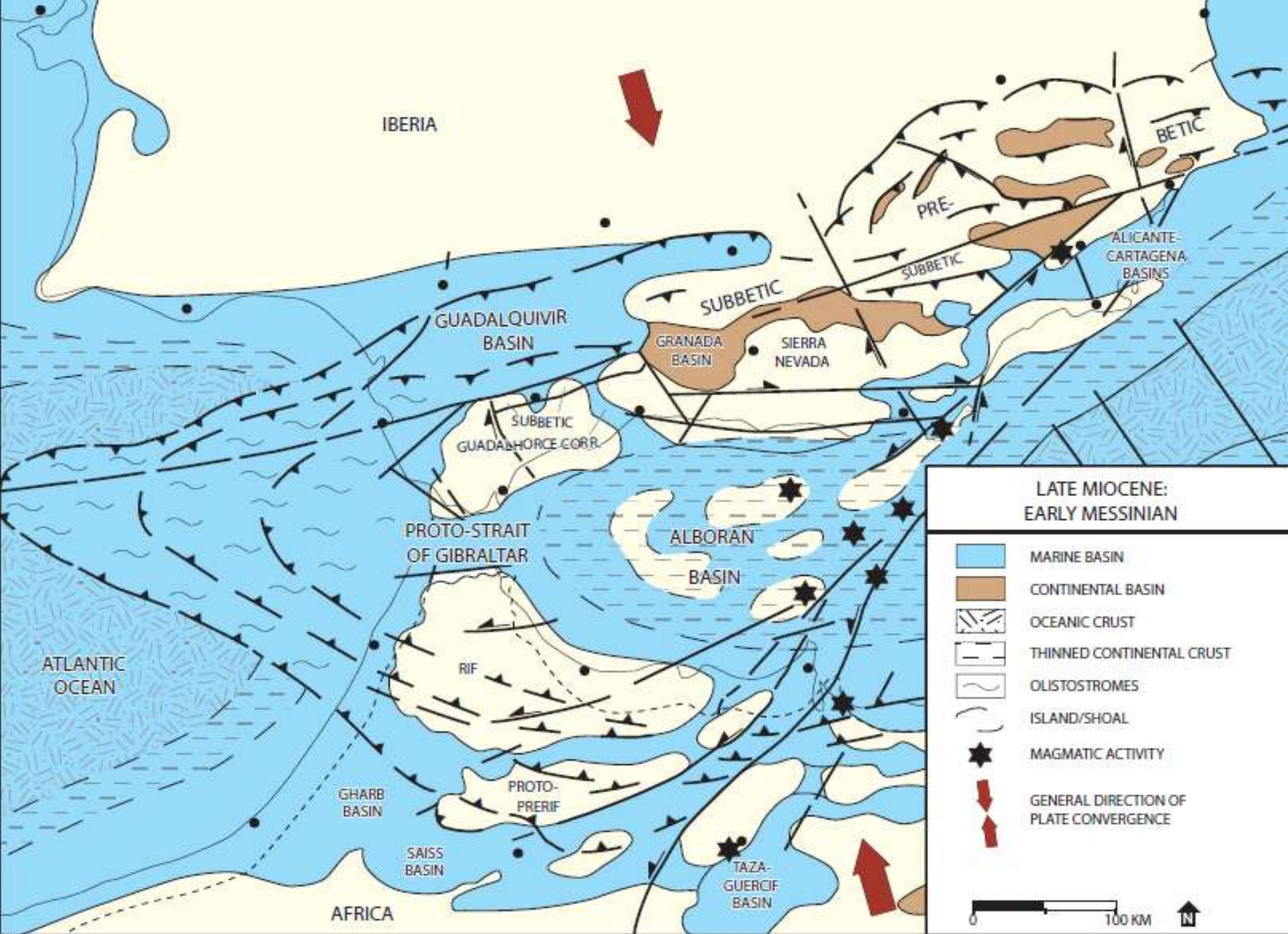
GHARB BASIN

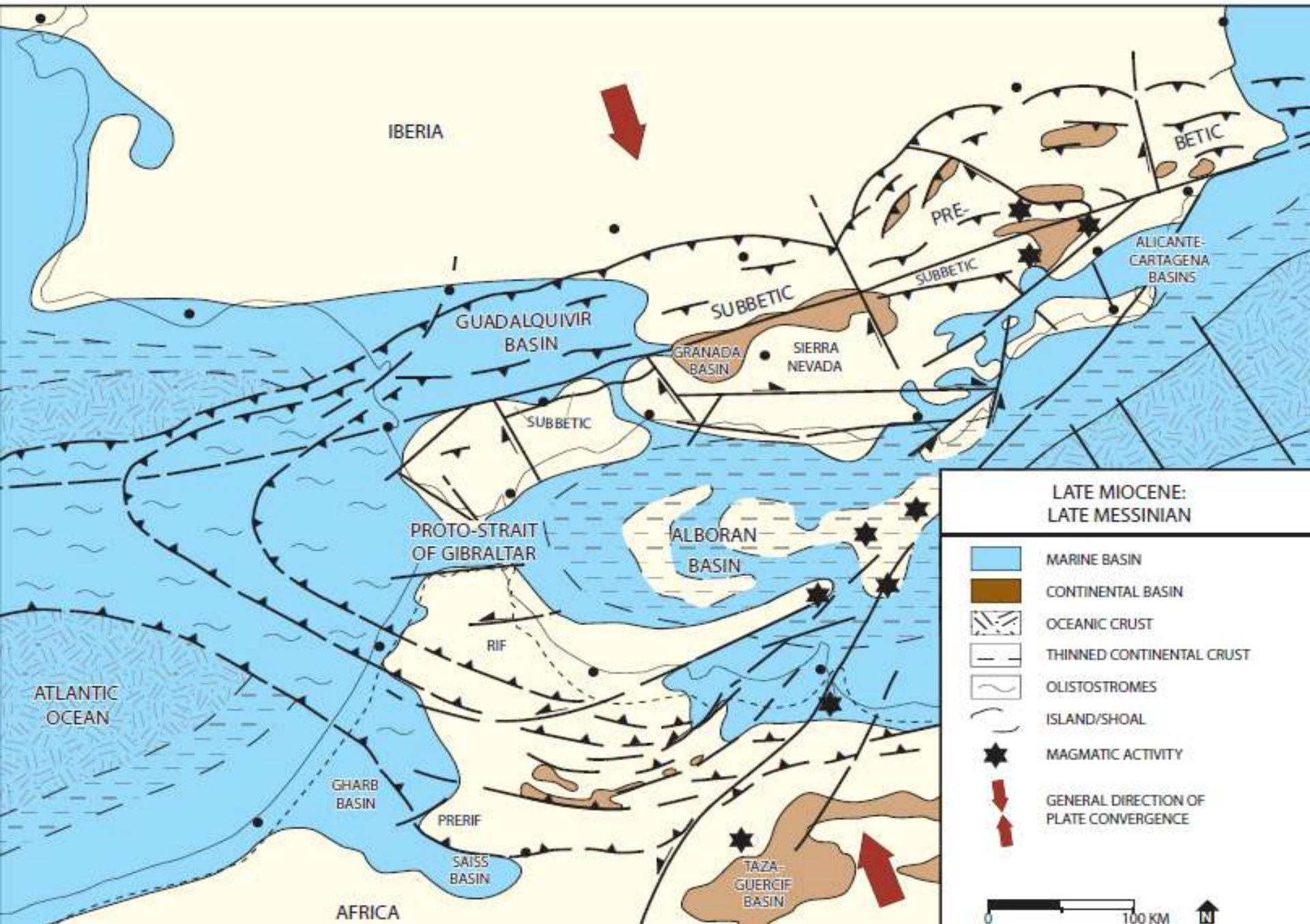
PROTO-SOUTH RIFIAN CORRIDOR

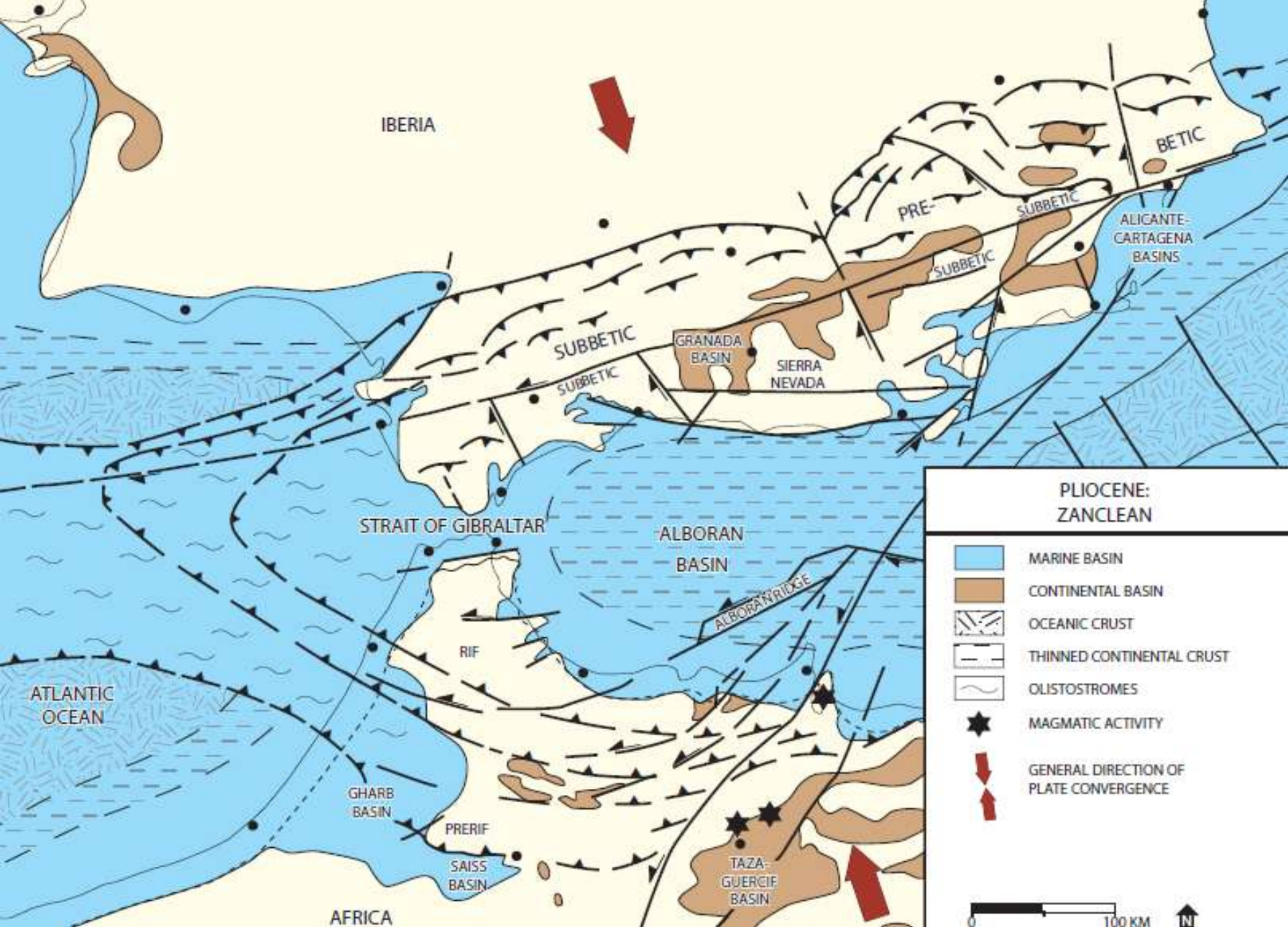
SAISS BASIN

AFRICA





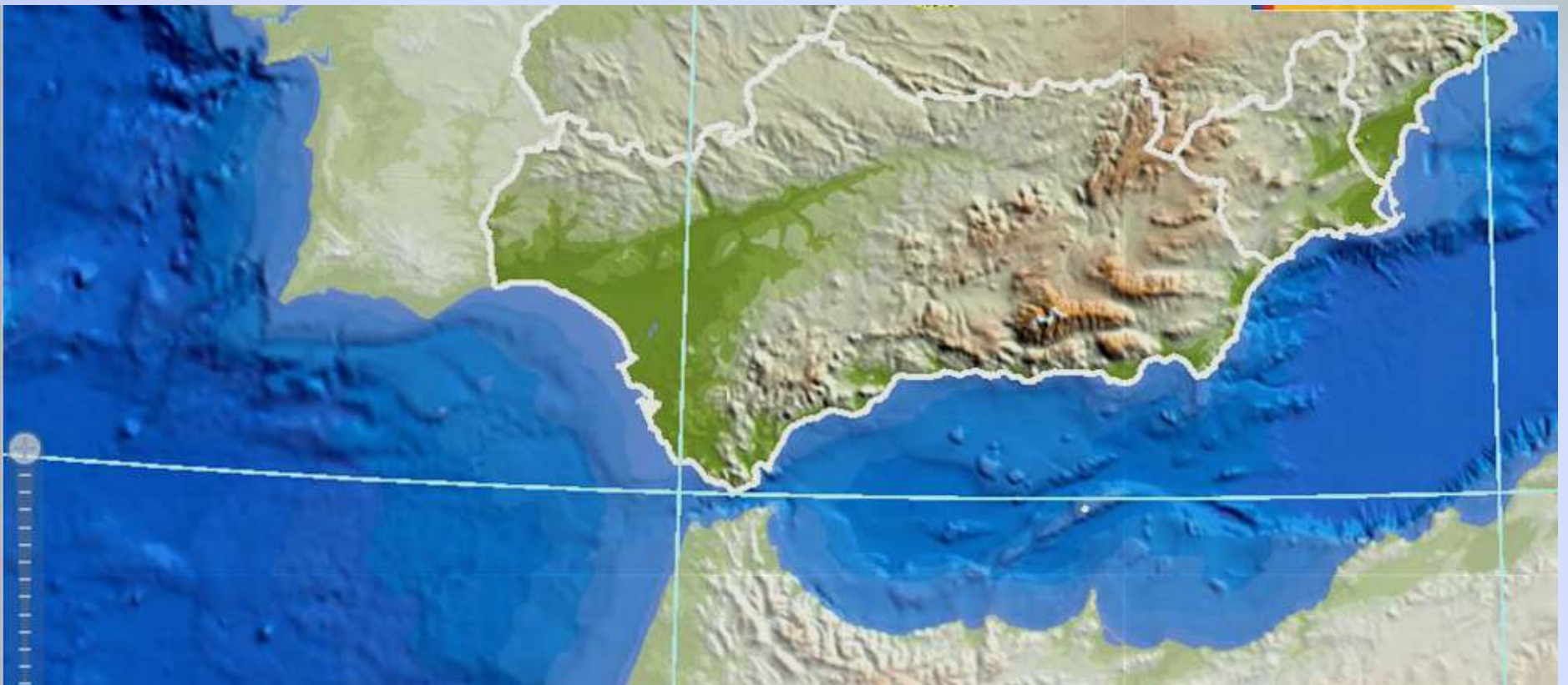




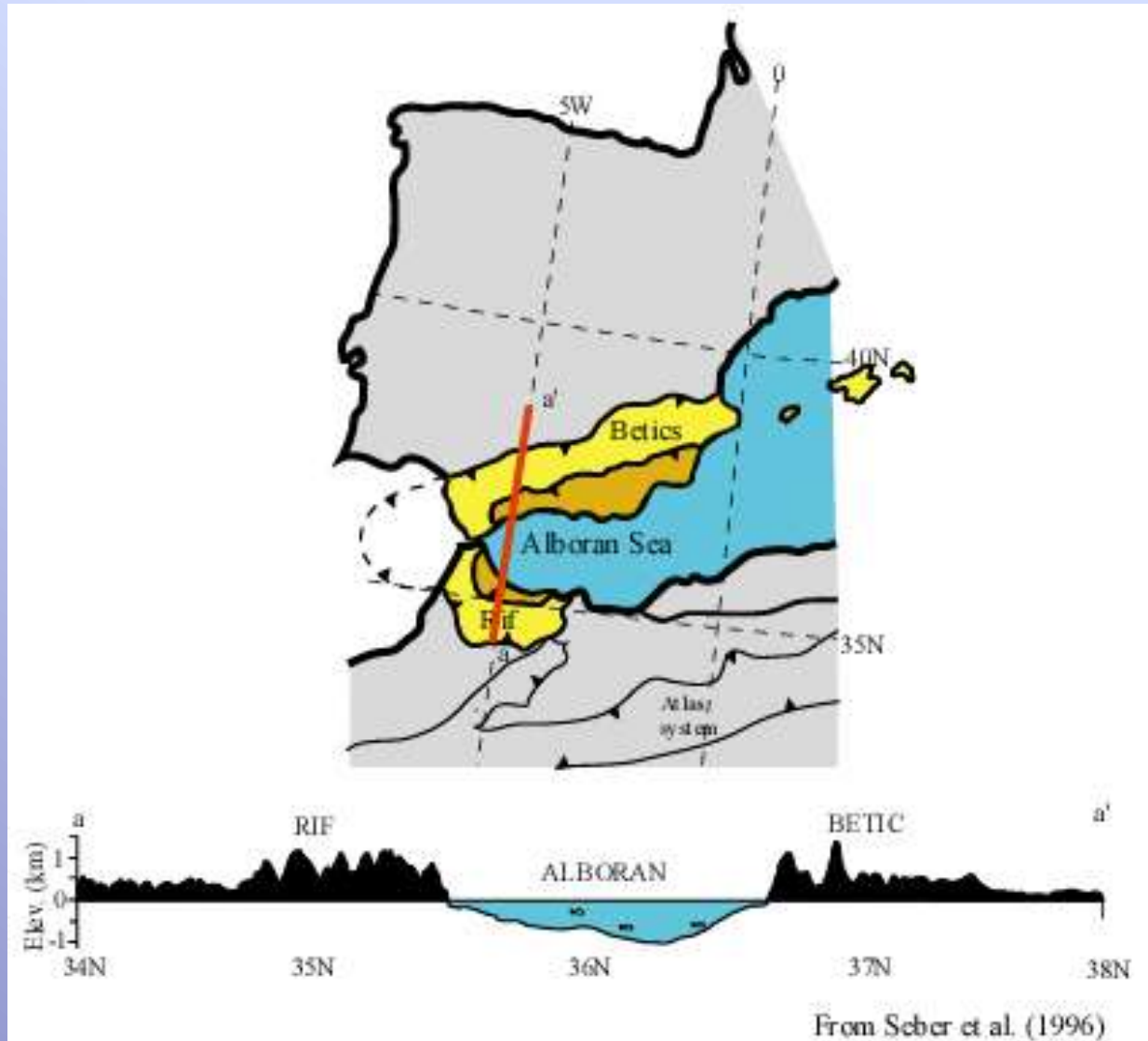
Wat hebben we zojuist gezien?

The Betic Cordillera, lying just between the Atlantic Ocean and the Mediterranean Sea, started to uplift in the Middle Miocene as result of the Alpine collision. First, only certain minor reliefs emerged as small, isolated islands. In Late Miocene, the paleogeography evolved into irregularly trending, interconnected, intermontane marine basins as the emerging reliefs expanded in size and fused together. Some of these basins opened directly to the Atlantic Ocean and others to the Mediterranean Sea (Braga et al., 2003). In the initial stages, numerous marine passages formed between the Atlantic-linked and the Mediterranean-linked basins. These seaways were, however, progressively confined and reduced in number during the uplifting until they finally disappeared.

Straat van Gibraltar



Structuur en topografie







Sierra Nevada



Marokkaanse Mediterrane kust



Rif Chefchaouen mountains



De Spaanse zeestraten in meer detail

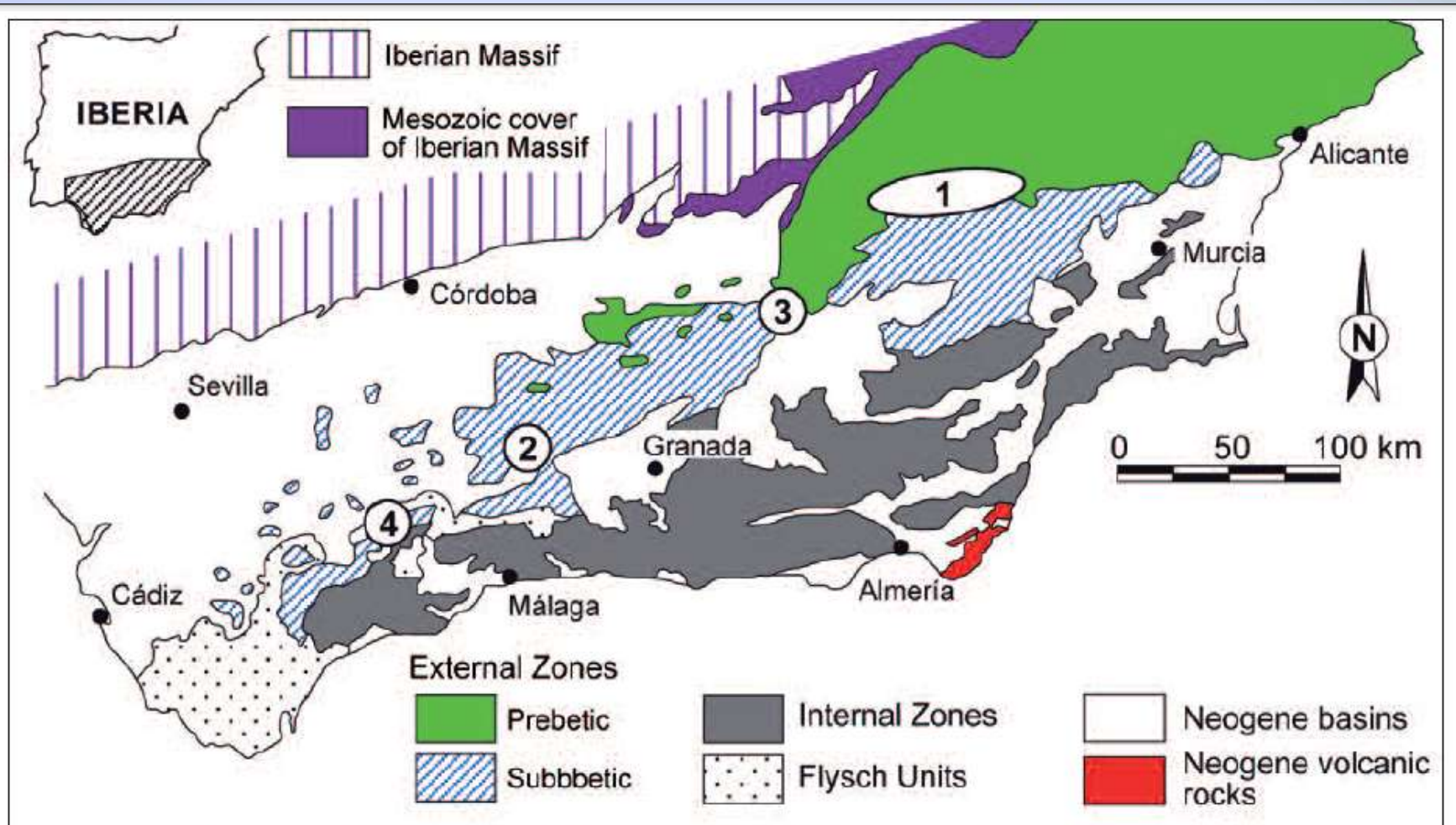


Fig. 1.- Simplified geological map of the Betic Cordillera. The position of the successive Betic straits is marked. 1: North-Betic Strait; 2: Zagra Strait; 3: Dehesas de Guadix Strait; 4: Guadalhorce Strait.

Crossbedding

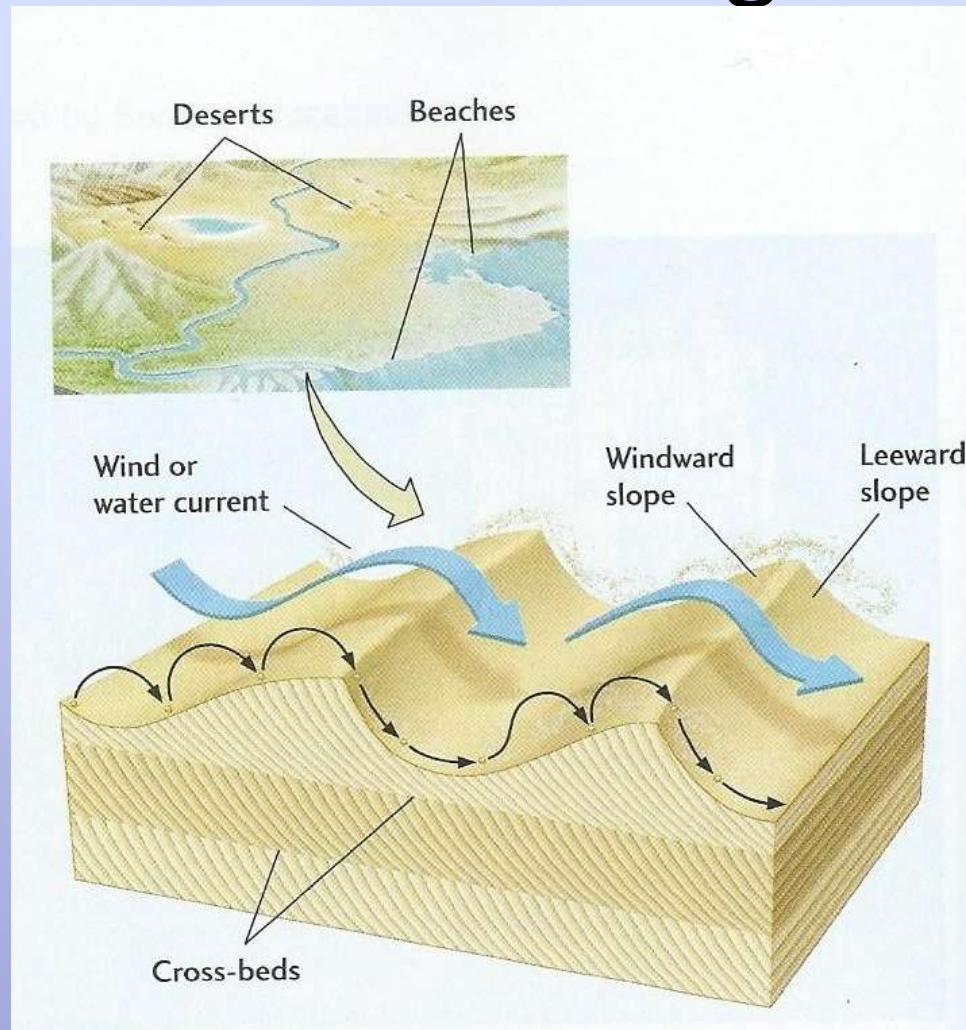
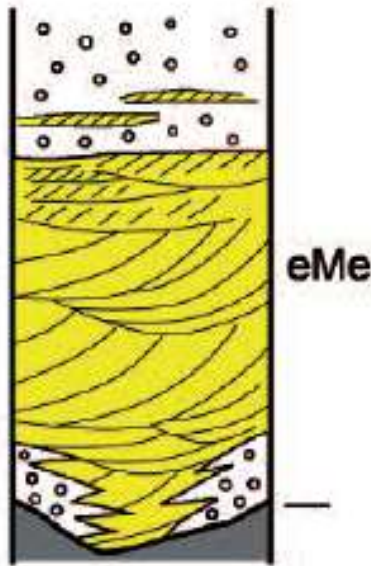


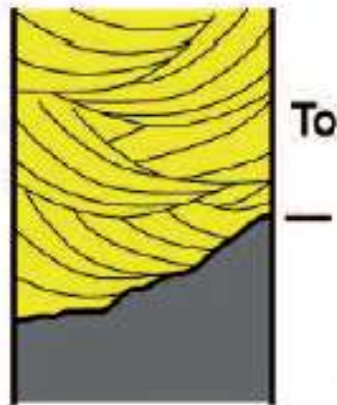
FIGURE 5.11 ■ Sediment particles transported down the steeper, downcurrent slope of a sand dune, sandbar, or ripple form cross-bedding.

De gelaagdheid van de zeestraten

Guadalhorce Strait



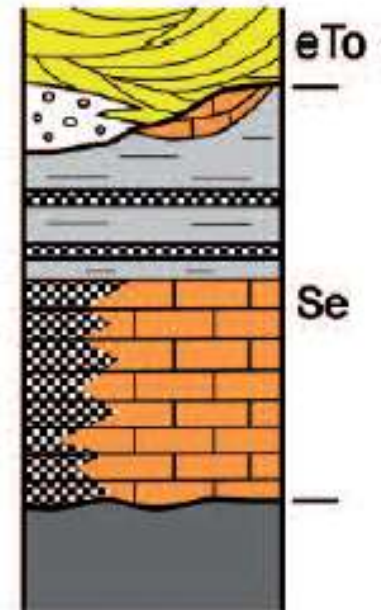
Zagra Strait



Dehesas de Guadix Strait



North Betic Strait



50 m

— unconformity

■ Middle Miocene and older Betic basement

■ continental deposits

■ marine conglomerate and sandstone

■ mixed bioclastic limestone and siliciclastic deposits with large-scale cross-bedding

■ temperate-water limestone

■ marine sandstone

■ marl

■ reef



North Betic Strait

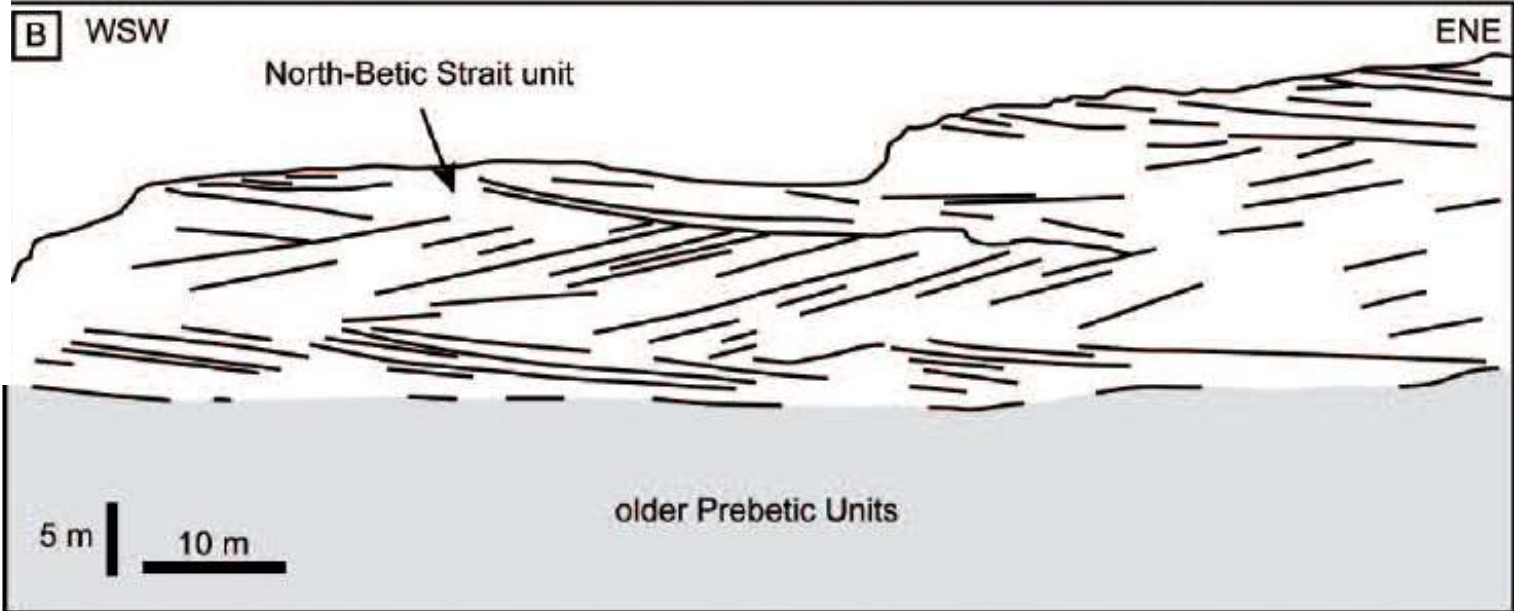
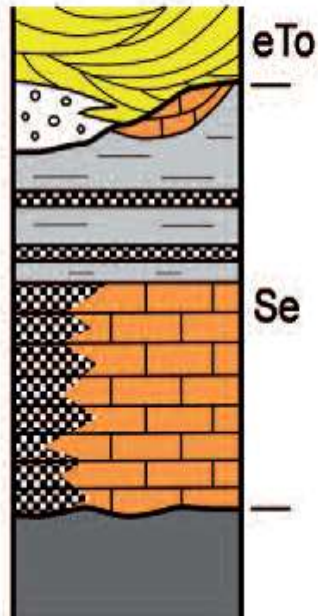


Fig. 3.- Composite structures exhibiting reverse, large-scale cross-bedding are common features in the North-Betic Strait deposits

(b)

Tortonian

Zagra Strait

DGS

Almería

Rifian straits

100 km





Fig. 4.- A-B: Huge, cross-bedding structures pointing to the SE. A collapsed dune, exhibiting strongly distorted and folded beds, can be seen at the bottom right of the picture. C: Close-up view of the area marked in A. Fuentes de Cesna (Zagra Strait).

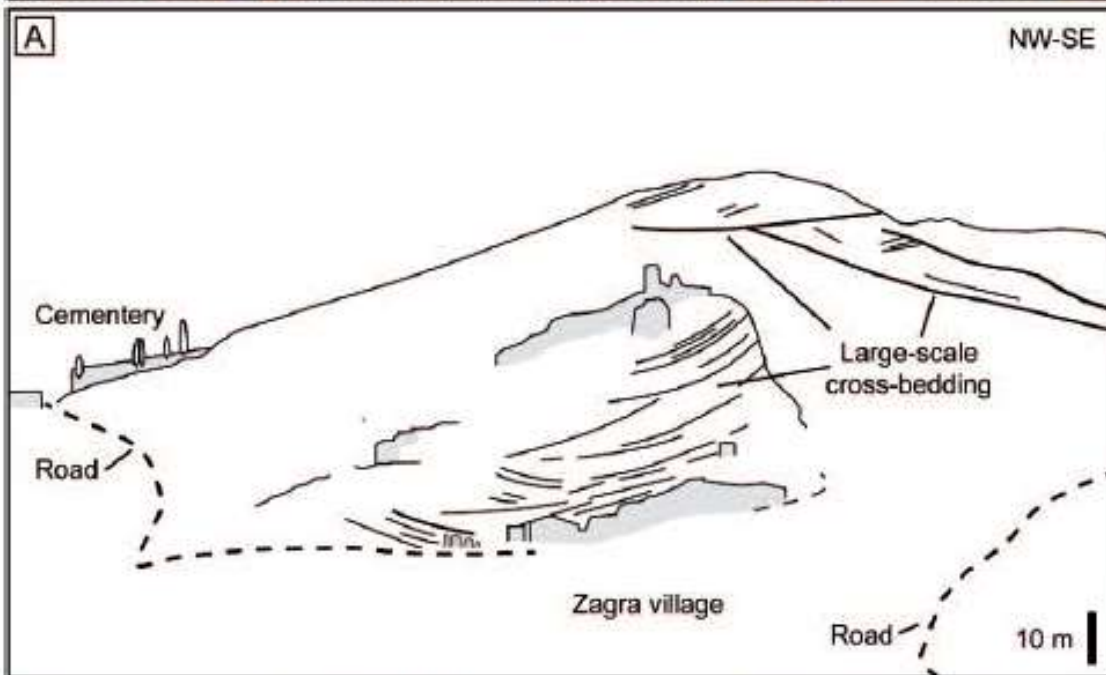


Fig. 5.- S-SE and N-NW dipping giant trough cross-bedding units are ubiquitous in the Zagra Strait.

A

NW-SE



(b)

Tortonian

Zagra Strait

DGS

Almería

Rifian straits

100 km



N-S



Fig. 6.- Unidirectional, giant cross-bedding is the most outstanding feature in the Dehesas de Guadix Strait. Man in the picture (encircled) is 1.85 m tall.

Dehesas de Guadix Strait



(c) earliest Messinian



Guadalhorce Corridor

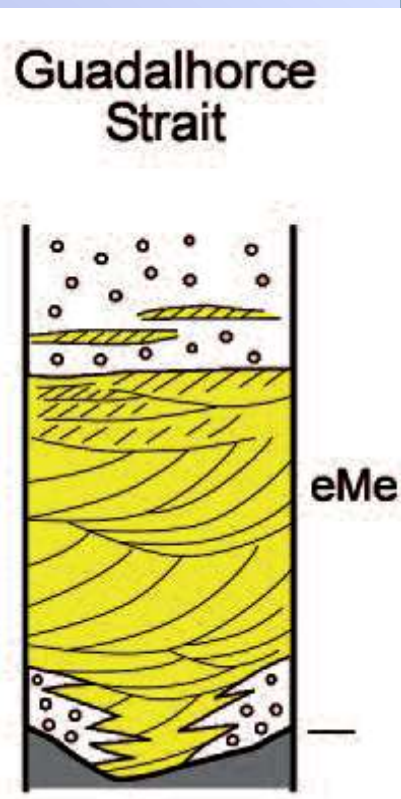


Fig. 7.- Aerial view of the Guadalhorce Corridor: Horizontal to gently dipping layers of the Messinian strait-infilling stand out in the centre of the picture. The Miocene beds are limited by Jurassic limestones (upper part of the picture) and by Paleozoic schists (bottom part of the picture).

An aerial photograph of the Guadalhorce Corridor, showing a complex mountainous landscape with various geological features. A compass rose in the top left corner indicates North. The wing of an airplane is visible in the top right corner. The text "Guadalhorce Corridor" is overlaid in the upper center of the image.

Guadalhorce Corridor

Fig. 7.- Aerial view of the Guadalhorce Corridor: Horizontal to gently dipping layers of the Messinian strait-infilling stand out in the centre of the picture. The Miocene beds are limited by Jurassic limestones (upper part of the picture) and by Paleozoic schists (bottom part of the picture).

Rifian corridors



Problemen Evaporieten MSC

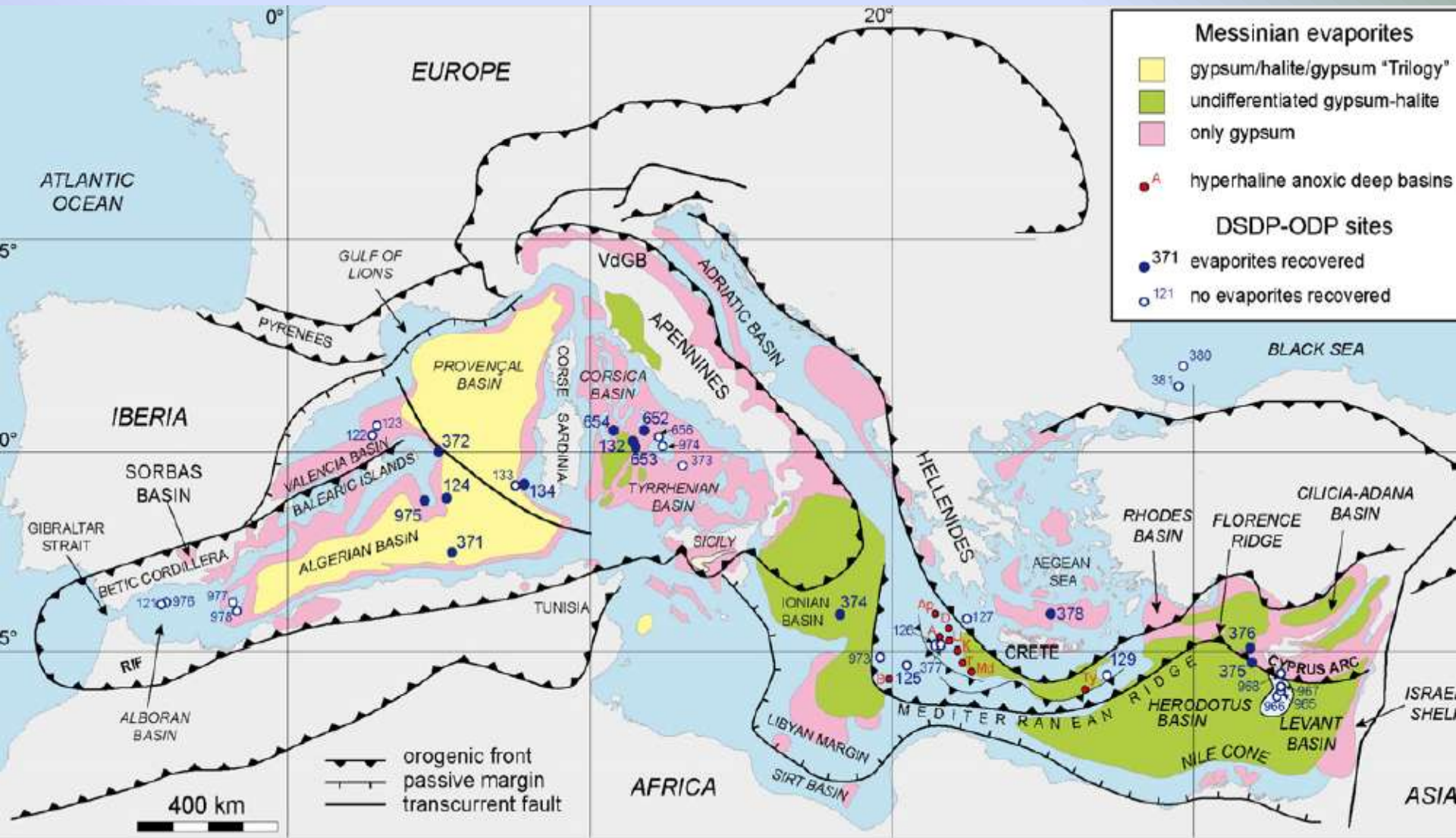
- Geen boringen door totale afzettingen in de diepzee, info alleen uit bovenste laag en ondiepere bekken (in zee of op het land)
- Er was geen adequaat stratigrafisch tijdsframe
- De afzettingen waren zo wie zo complex
- Wel was er een spectaculaire cycliciteit in de gesteente afzettingen van de evaporieten te ontdekken
- Daardoor geschikt voor astro-chronological tuning en gedetailleerde bestudering

Gevolgen van een geleidelijke afsluiting? (de aanloop)

- 7,15 Ma vermindering van de diepzee circulatie
- Daarmee samenhangend een toenemende afzetting van diatomeeën rijke sedimenten de Tripolo formatie op o.a. Sicilië (maar over hele MZ)
- Ook opaal rijke afzettingen in Zuid-Spanje en door de hele MZ (7,15-6,7 MA)
- 6,7 De diversiteit van kalkhoudende plankton vermindert sterk (door zouter oppervlakte water)
- Tussen 6.3 en 5.97 calcië-,dolomiet- en aragonietvorming

Het begin: afzetting van de eerste gipslagen

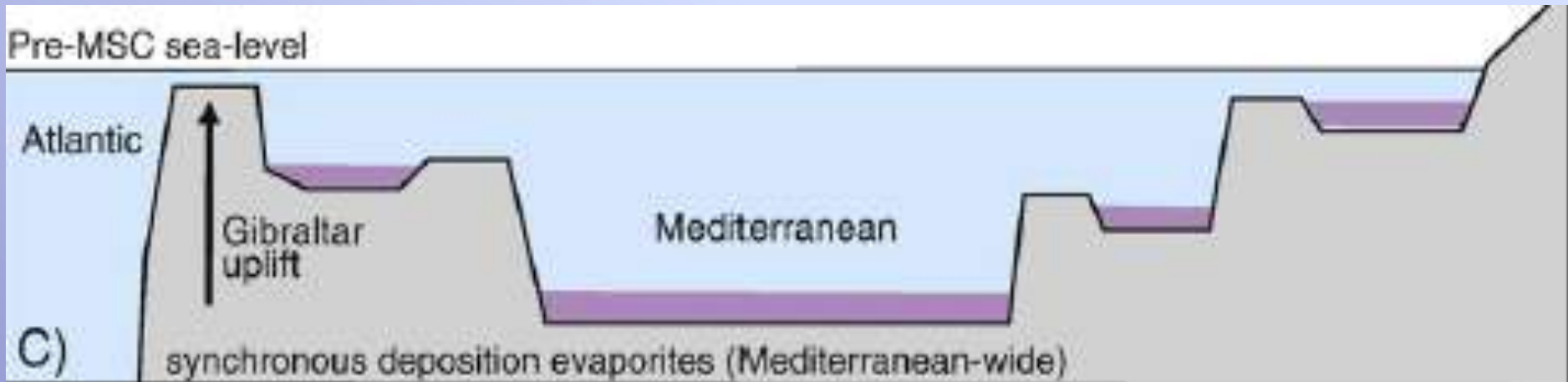
- Waar komen die afzettingen voor ?
- Hoe ontstaan?
- Hoe zien ze er uit?

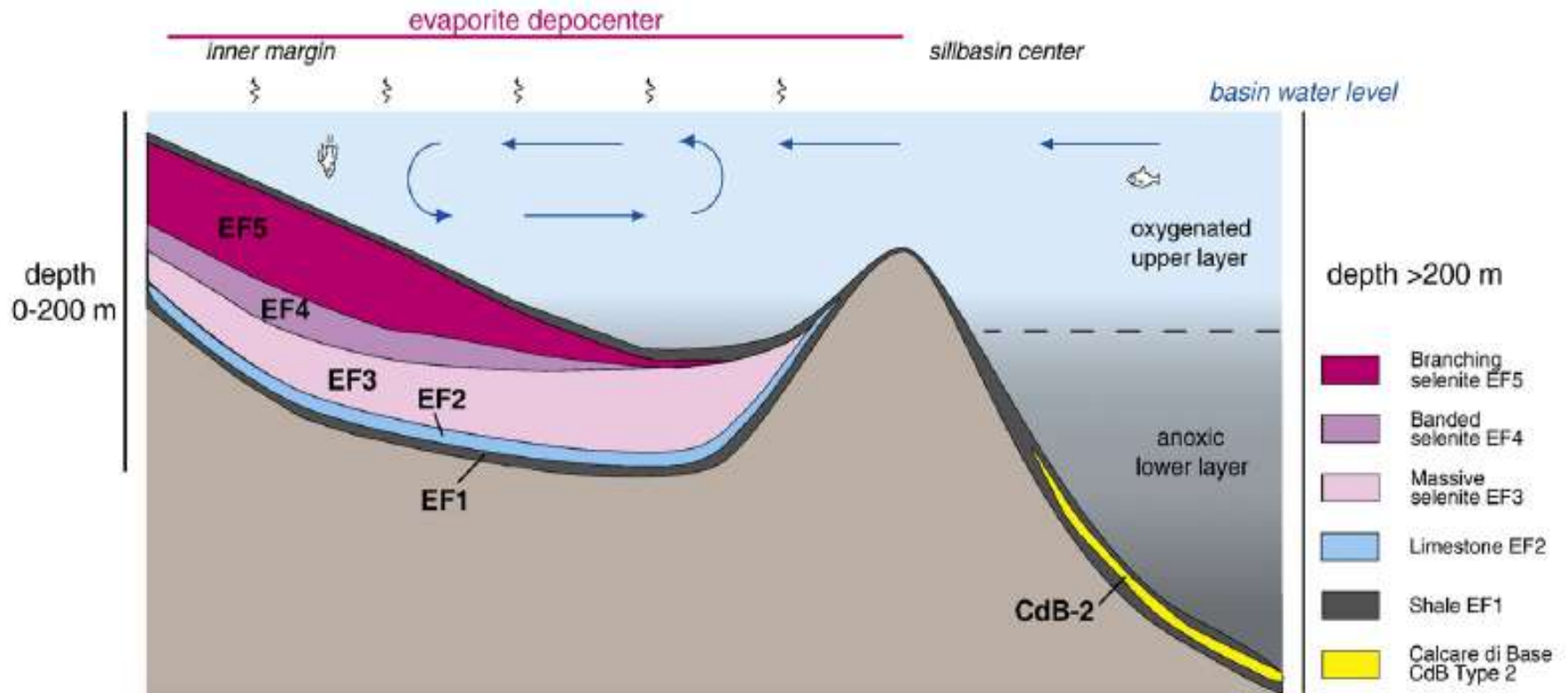
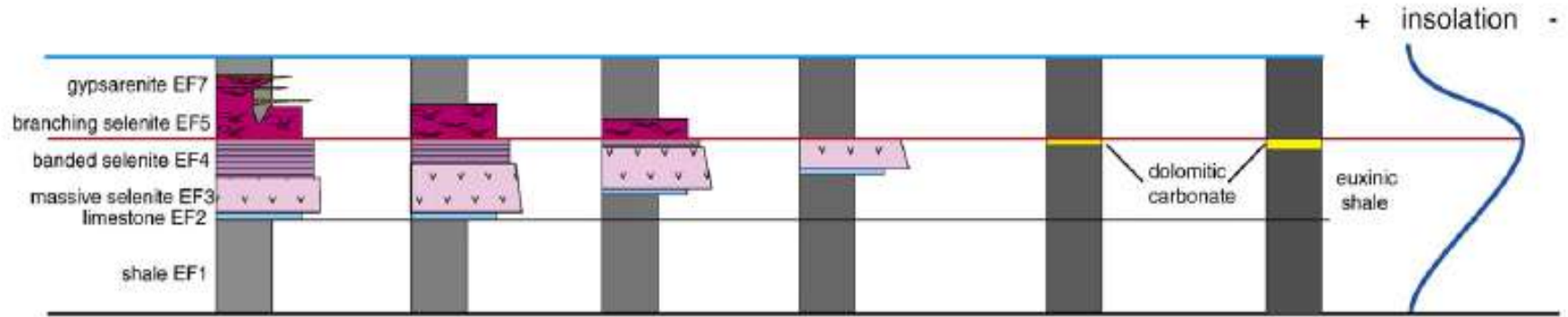


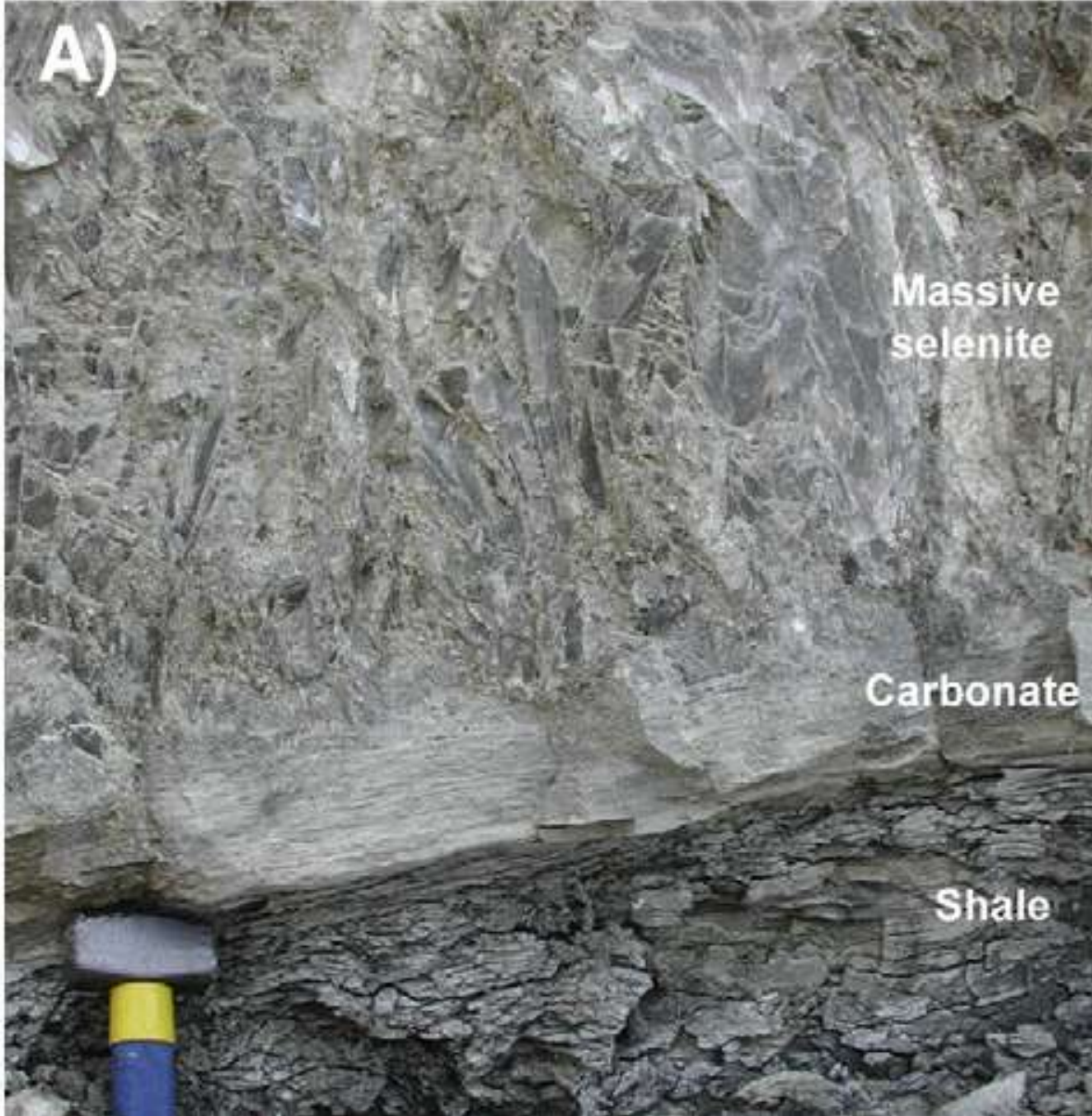
Omstandigheden tijdens vorming PLG's

- Relatief homogeen water dat zijn oorsprong vond uit de A.O.
- Er was een beperkte uitstroom van water uit de MZ naar de A.O.
- De instroom van rivierwater was ook belangrijk, significant
- Er was dus verbinding met de A.O.
- Relatief lage drempel tussen MZ en AO
- Continentale invloed en invloed van klimaat

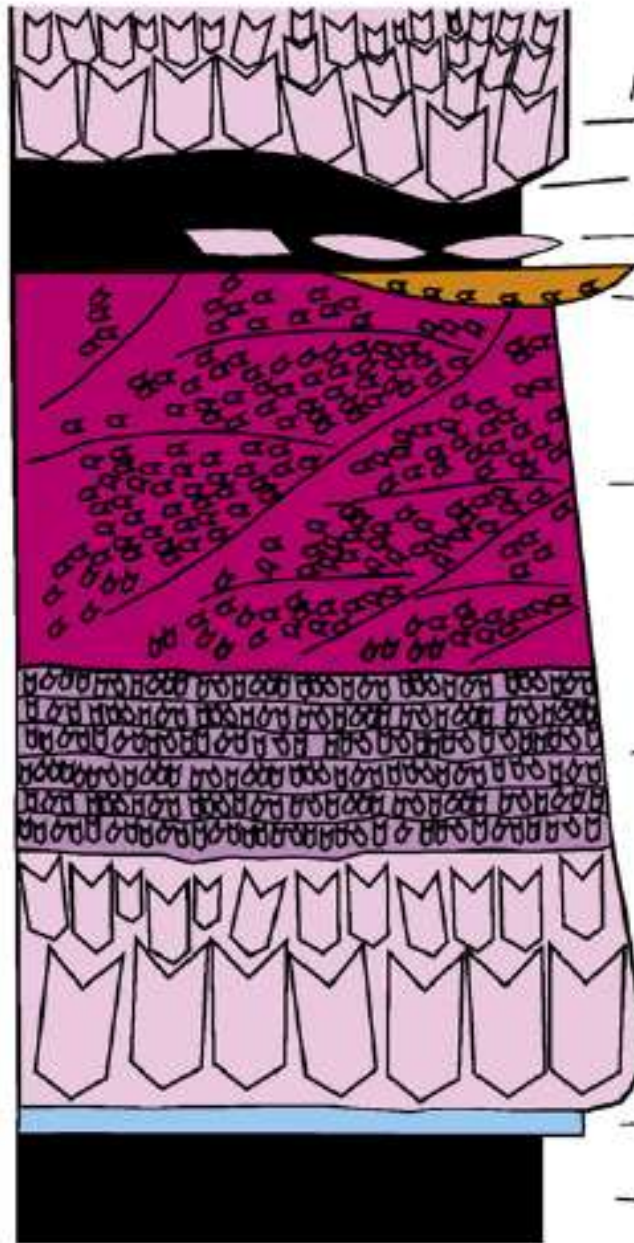
Zeespiegelniveau MZ=AO drempel wordt hoger







Primary Lower Gypsum facies association



Nucleation cone
("cavolo")
"Mammellone"

This work

- EF6 - Displacive selenite
- EF8 - Gypsarenite
- EF7 - Gypsrudite

- EF5 Branching selenite
(supercones in Sorbas)

- EF4 - Banded selenite

- EF3 - Massive selenite
(giant selenite > 0.3 m)

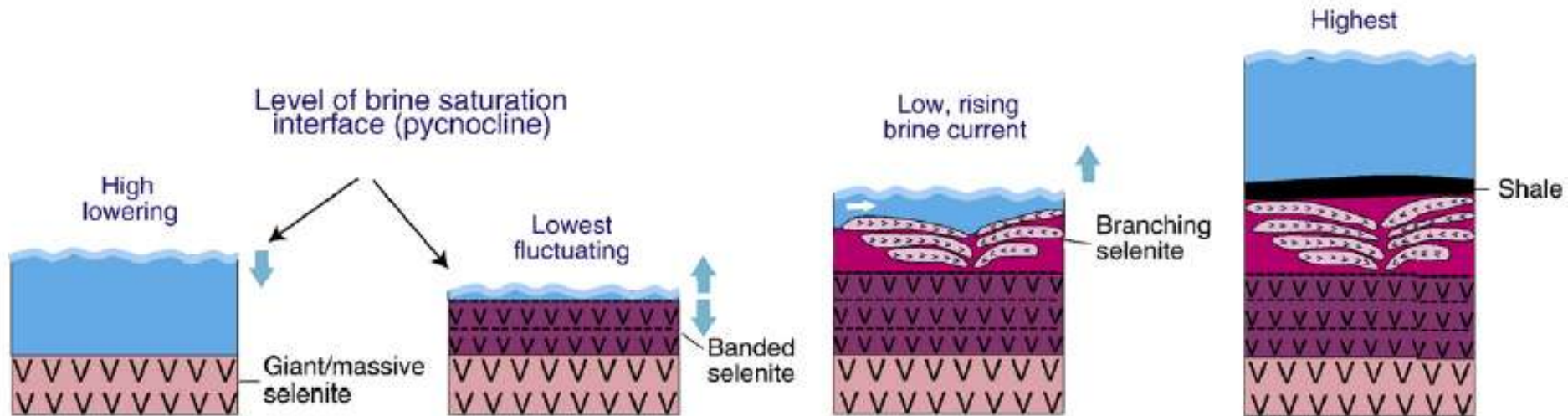
- EF2 - Limestone: massive,
laminated (stromatolite), breccia

- EF1 - Bituminous shale

Climatic precession cycle

ARID

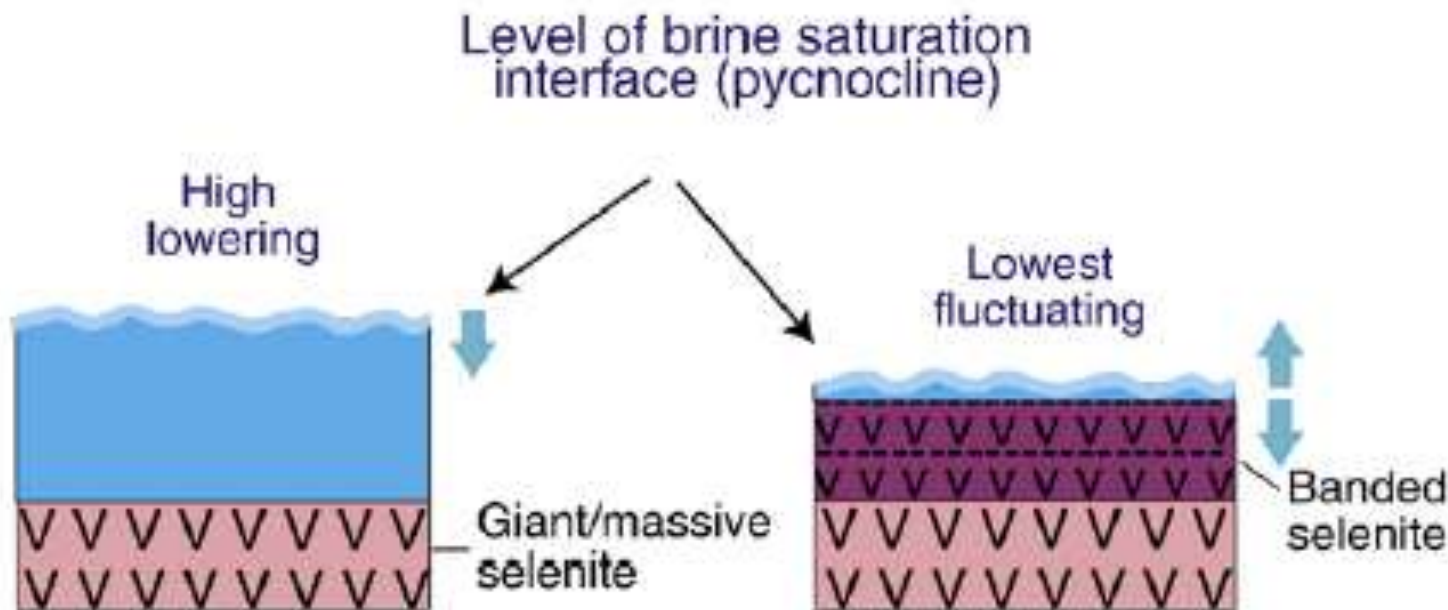
HUMID



Ontstaan Giant- of Massive Selenite

Climatic precession cycle

ARID





Competitive growth:
only vertically-oriented
nuclei have free space

Cyanobacteria filaments
trapped on the reentrant
angle of the twin

Swallow-tail
twinning plane

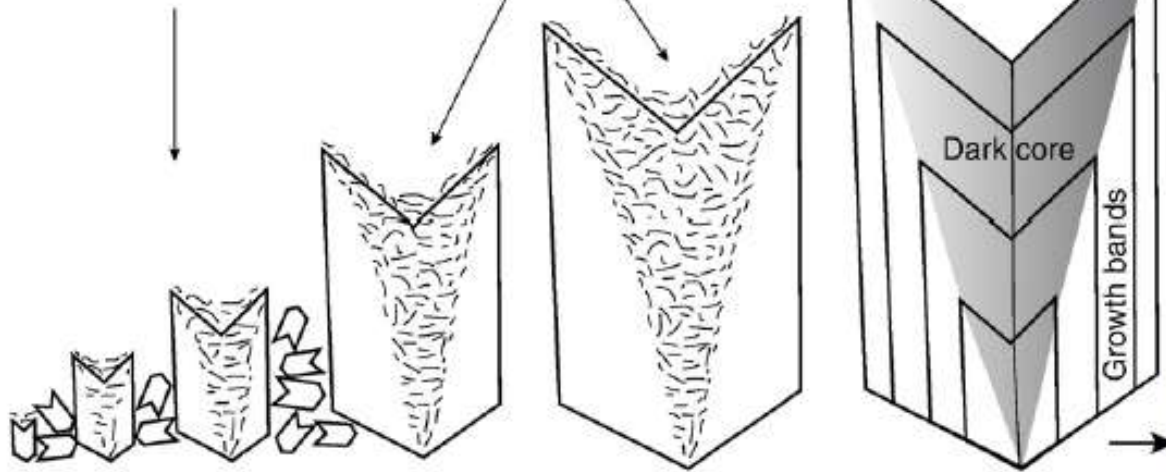
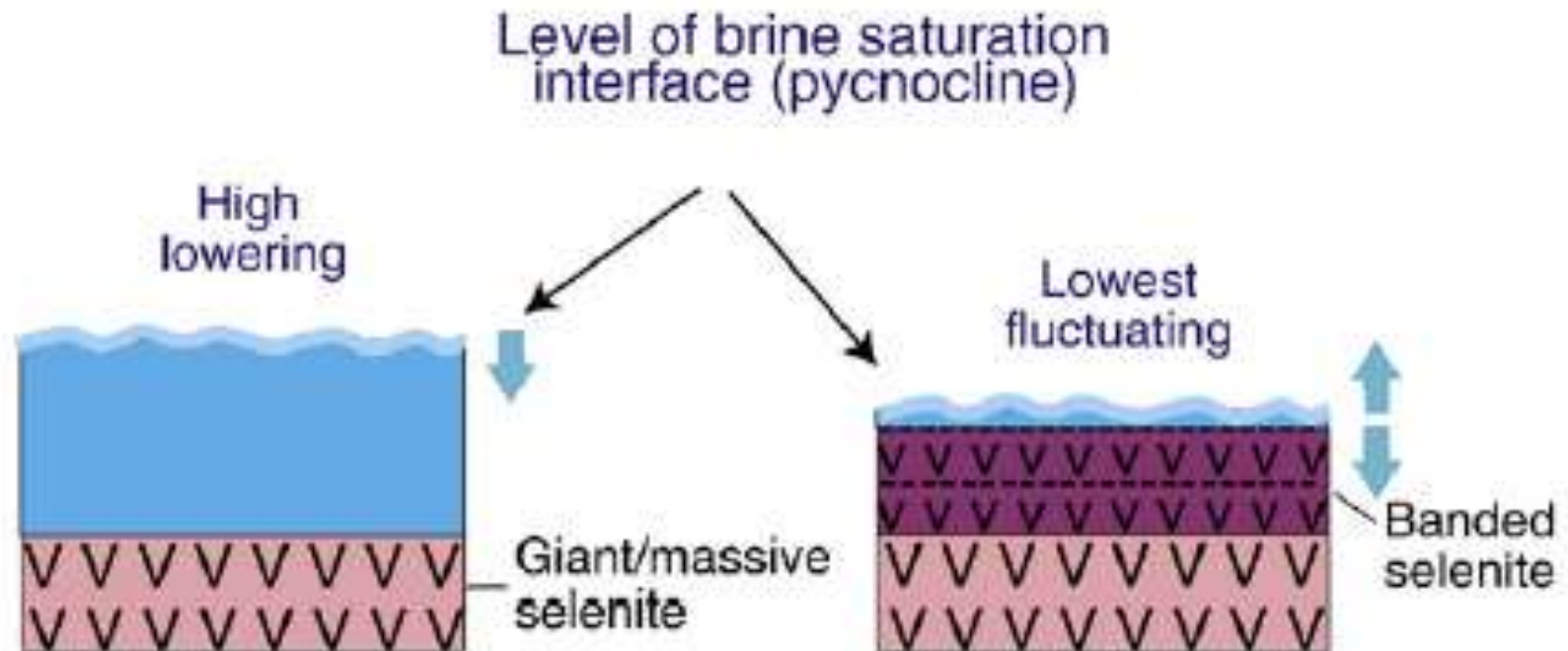


Fig. 8. Diagram showing the competitive growth of selenite determining the typical vertical orientation of the crystals (Mottura's rule) and the trapping mechanism of cyanobacteria filaments on the re-entrant angles of the twins producing the dark triangular core of the swallow-tail crystals.

Ontstaan banded Selenite

Climatic precession cycle

ARID





Overgang van Massive – naar banded selenite

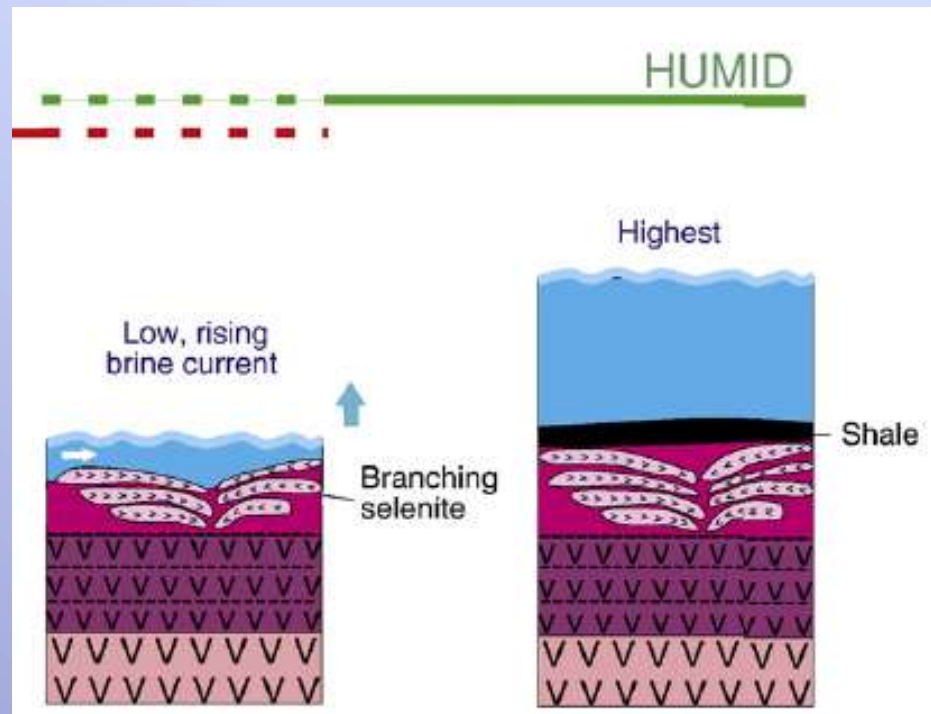


c)



2 cm

Ontstaan Branching selenite



Branching selenite



Geraadpleegde literatuur 1

The Messinian Salinity Crisis: Past and future of a great challenge for marine sciences



Marco Roveri^{a,b,*}, Rachel Flecker^c, Wout Krijgsman^d, Johanna Lofi^e, Stefano Lugli^f, Vinicio Manzi^{a,b}, Francisco J. Sierro^g, Adele Bertini^h, Angelo Camerlenghiⁱ, Gert De Lange^j, Rob Govers^j, Frits J. Hilgen^j, Christian Hübscher^k, Paul Th. Meijer^j, Marius Stoica^l

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Geraadpleegde literatuur 2

The Messinian Salinity Crisis: Past and future of a great challenge for marine sciences



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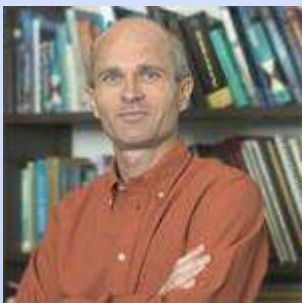
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2 Marco Roveri



3 prof. dr. Wout Krijgsman



5 Vinicio Manzi



6 D. Garcia-Castellanos

Einde